Effects of Surface Layer Physics Schemes on the Simulated Intensity and Structure of Typhoon Rai (2021)



Thi-Huyen Hoang, Ching-Yuang Huang, Thi-Chinh Nguyen Department of Atmospheric Sciences, National Central University

Sep 5th, 2024





Introduction

- Numerical simulation depends on many factors: initial condition, initial time, physical processes ...
- Influence of microphysics scheme, cumulus parameterization, and PBL parameterization on TC track and intensity (Kanase et al. 2015; Miglietta et al. 2015)
- But few studies have discussed the impact of surface layer (SL) physics schemes on TC simulation
- SL physics scheme: determine the exchange of momentum, heat, and moisture (Skamarock et al., 2021)
- Crucial for the formation, development, and maintenance of tropical cyclones (Byers 1944; Riehl 1954; Malkus and Riehl 1960; Emanuel 1986; Gao et al. 2019)







Typhoon Rai (2021)



Powerful and destructive TC hit the Philippines

Typhoon Rai (Odette)

Peak intensity: 105 kt (54 m/s), 915 hPa off the Philipines coast (06 UTC on December 16)

Extreme rain, strong wind, and storm surge: over 400 lives, over \$1.02 billion U.S. dollars in damages



The big tropical cyclone impact highlighted the need for improved forecasting of track and intensity

Experimental Design

	Rai (2021)				
30 ⁰ N -	d01 d02 d03				
20 ⁰ N -					
10°N -					
0° -					
10 ⁰ S -					

120⁰E

90⁰E

105⁰E

Model	WRF-ARW model (version 4.5.1)		
Domains	Domain 1: 18 km (fixed) Domain 2: 6 km (moving) Domain 3: 2 km (moving)		
Vertical levels	51 levels		
Time runs	96 hours		
Data	NCEP FNL		

Cumulus prameterization	KF
Microphysics	Lin
PBL parameterization	MYNN
Land surface model	Noah
Shortwave radiation	Dudhia
Longwave radiation	RRTM

SL physics schemes: MO, MM5, MYNN

135⁰E

parameterization of surface fluxes plays key role in the simulation

150⁰E

165⁰E





- ✓ Simulated Track and Intensity
- ✓ Characteristics of Surface Layer
- ✓ Characteristics of Planetary Boundary Layer
- ✓ Typhoon Structure
- ✓ Solution of Sawyer-Eliassen Equation

Simulated Track and Intensity



Assessment statistics of simulated maximum wind speed							
	R	BIAS (m s ⁻¹)	RMSE (m s ⁻¹)				
CTL	0.995	-0.412	0.849				
MM5	0.930	-4.270	5.378				
MYNN	0.932	-5.359	5.745				

00 UTC 15 Dec to 06 UTC 16 Dec 2021

Compare with JMA R: correlation coefficient BIAS: the bias RMSE: Root mean square error

SL physics schemes barely affect TC track but strongly affect TC intensity

CTL good produce the Typhoon Rai intensity

 \rightarrow TC intensity is **sensitive to** the selection of different SL physics schemes

Characteristics of Surface Layer





Characteristics of Boundary Layer



Black line: height of maximum wind speed

- Locations of maximum V_t and V_r are in different regions
 - KinematicboundarylayerheightdecreaseswithdecreasingradiustothecenterVVV
- $h V_r > h V_t$
- **CTL**: larger circulation and PBL height

Black line: height where radial wind speed is 10% of peak inflow



MM5

CTL



Typhoon Structure

Vector: radial and vertical wind Green line: radius of maximum wind speed



CTL: stronger inflow, upward, and diabatic heating

Solution of Sawyer-Eliassen Equation

Horizontal advection

Vector: radial and vertical wind



- Turbulent friction: important in the low-level
- Diabatic heating: important in the low-level and upper level

Solution of Sawyer-Eliassen Equation

Vertical advection

Vector: radial and vertical wind



Diabatic heating dominant the vertical advection

Conclusions

- SL layer physics schemes: minimal affect TC track but strongly affect TC intensity.
- The highest correlation coefficient and smallest bias and root mean square error between the best track is in CTL compared to MM5 and MYNN
 The CTL produces good Typhoon Rai intensity → good option for SL physics scheme.
- Relatively higher surface wind speed, friction velocity, ratio of exchange coefficient, surface fluxes of heat and moisture are found in CTL, followed by MM5 and MYNN
- TC structures: primary and secondary circulation, potential temperature, boundary layer heights, warm-core strength and height are substantially impacted by surface layer physics schemes.
- Using the SE equation to analyze the contributing components, diabatic heating and turbulent friction play a major role in the secondary circulation.

Key findings <</p>

References

- Byers, H. R., 1944: General Meteorology. McGraw-Hill, 645 pp.
- Emanuel, K.A., 1986. An air-sea interaction theory for tropical cyclones. Part I: Steady-state maintenance. *Journal of Atmospheric Sciences*, **43**(6), pp.585-605.
- Gao, S., Jia, S., Wan, Y., Li, T., Zhai, S. and Shen, X., 2019. The role of latent heat flux in tropical cyclogenesis over the western North Pacific: Comparison of developing versus non-developing disturbances. *Journal of Marine Science and Engineering*, 7(2), p.28.
- Kanase, R.D. and Salvekar, P.S., 2015. Impact of physical parameterization schemes on track and intensity of severe cyclonic storms in Bay of Bengal. *Meteorology and Atmospheric Physics*, **127**, pp.537-559.
- Malkus, J.S. and Riehl, H., 1960. On the dynamics and energy transformations in steady-state hurricanes. *Tellus*, **12**(1), pp.1-20.
- Miglietta, M.M., Mastrangelo, D. and Conte, D., 2015. Influence of physics parameterization schemes on the simulation of a tropical-like cyclone in the Mediterranean Sea. *Atmospheric Research*, **153**, pp.360-375.
- Skamarock, Klemp, J. B., Dudhia, J., Gill, D. O., Liu, Z., Berner, J., Wang, W., Powers, J. G., Duda, M. G., Barker, D., and Huang, X. -Y, 2021. A Description of the Advanced Research WRF Model Version 4.3. *No. NCAR/TN-556+STR*.
- Riehl, H., 1954: Tropical Meteorology. McGraw-Hill, 392 pp.



Sawyer-Eliassen Equation

Sawyer–Eliassen (SE) equation: derived as a linear partial differential equation applied to investigate the responses of a primary mean vortex, different force sources, and the intensification of the TC (Nguyen and Huang, 2023)

$$\mathbf{S} = -\frac{\partial}{\partial z} (\chi \xi \mathbf{G}_{\boldsymbol{v}}) + \frac{\partial}{\partial r} [\mathbf{g} \chi^2 (1 + \epsilon q_{\boldsymbol{v}}) \mathbf{H}_{\boldsymbol{\theta}}] + \frac{\partial}{\partial z} [C \chi^2 (1 + \epsilon q_{\boldsymbol{v}}) \mathbf{H}_{\boldsymbol{\theta}}] - \frac{\partial}{\partial z} (\chi \xi \dot{\mathbf{V}}) + \frac{\partial}{\partial t} \left(\frac{\partial \chi \dot{\mathbf{U}}}{\partial z}\right) - \frac{\partial}{\partial t} \left(\frac{\partial \chi \dot{\mathbf{W}}}{\partial r}\right)$$

(momentum)



- (1) Eddy transport
- (2) Turbulent momentum diffusion
- (3) Eddy temperature
- (4) Turbulent heat diffusion
- (5) Diabatic heating

(thermodynamic heat)

$$H_{\theta} = -u'\frac{\partial\theta'}{\partial r} - v'\frac{\partial\theta'}{r\partial\lambda} - w'\frac{\partial\theta'}{\partial z} + \bar{F}_{\theta} + \frac{1}{\bar{\Pi}}\bar{Q}$$
(3)
(4)
(5)

Residual terms \dot{V} , \dot{U} , \dot{W} and (4) are not significant compared to others,

are not shown

Radar Reflectivity



- CTL more symmetric than MM5 and MYNN, strongest and widest in CTL
- 42h: highly concentrated in the southern
- 54h: south-eastern

• 66h: south to south-western

12h Accumulated Precipitation



- Rainfall: concentrate along the path of TC
- First and second stage: CTL is strongest and widest
- Third stage: MM5 and MYNN are higher

Azimuthal-mean Angular Momentum $M = r\bar{v}$

Vector: radial and vertical wind Shaded: azimuthal-mean angular momentum Green line: radius of maximum wind speed



The AM decreases inward and upward with the TC center

CTL > MM5

21