

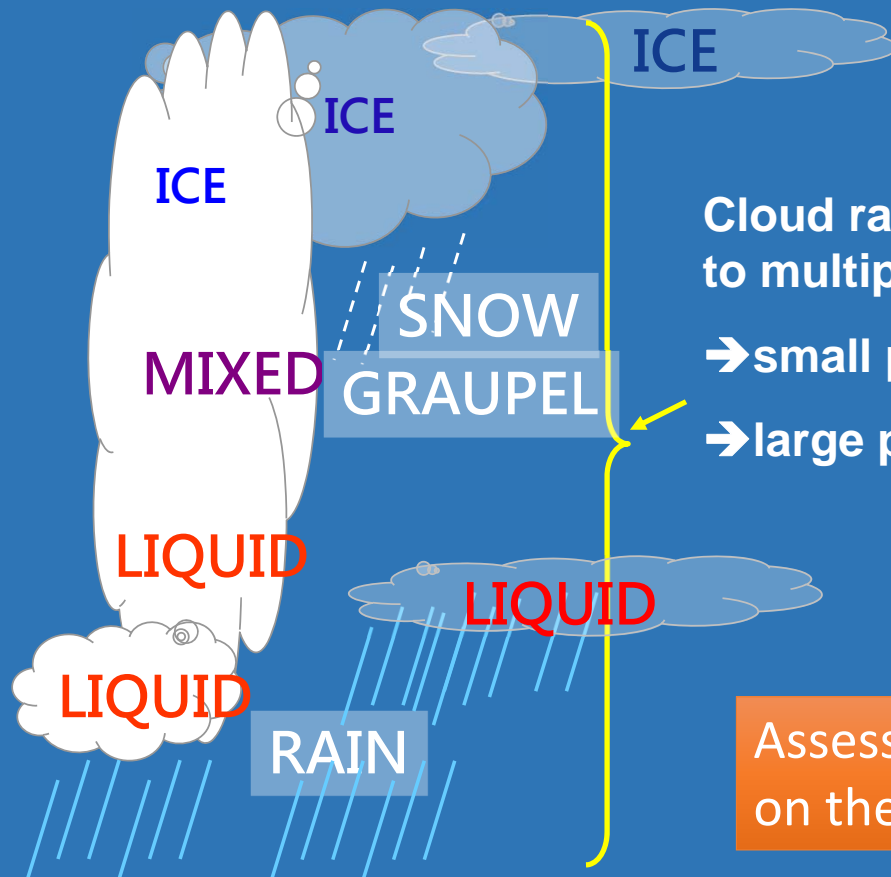
# WSM5 微物理方案 於CWB-GFS的評估

汪鳳如 陳建河

中央氣象局

- Motivation
- Description of the schemes
- Result

# motivation



Cloud radiation effect are sensitive to multiple particle types:

- small particle: cloud ice, liquid
- large particle: rain, snow, graupel

Assessment of the WSM5 MPS on the cloud radiation process

*From Jui-Lin Li*

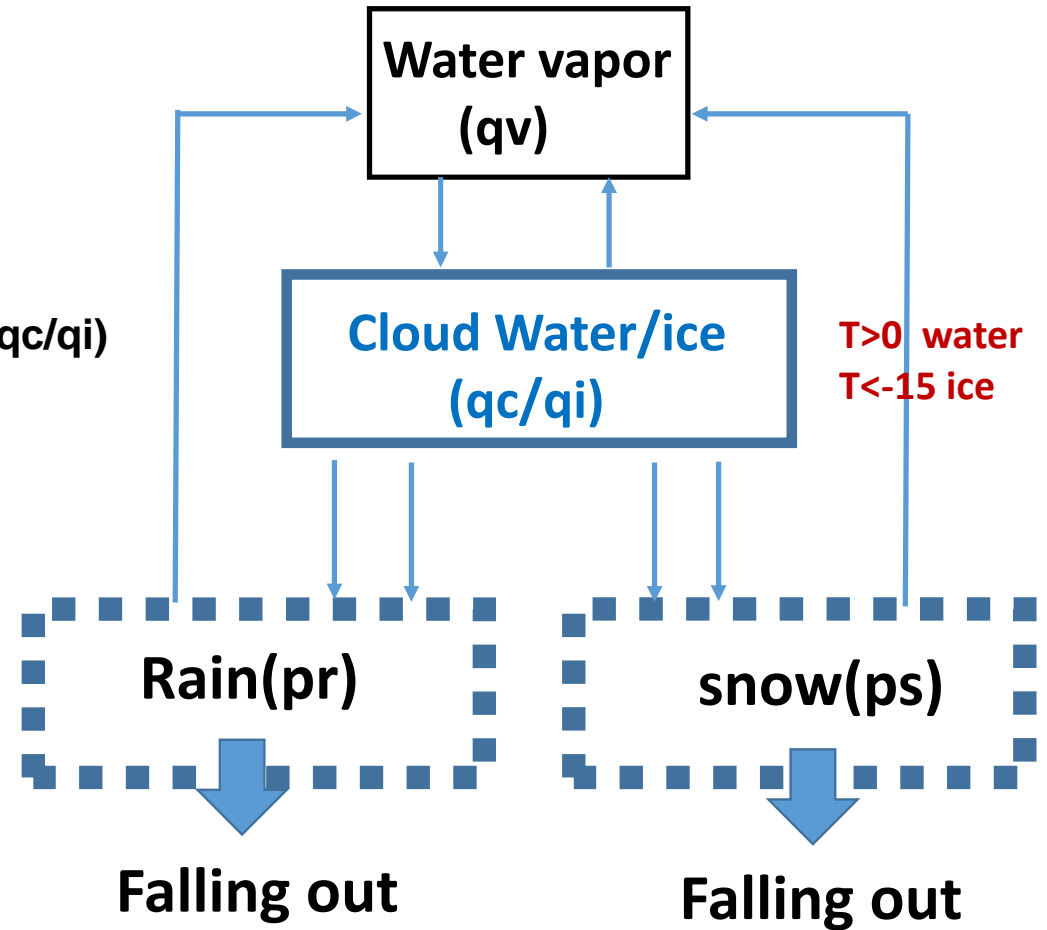
# CWB-GFS網格尺度降水方案

Explicit prognostic cloud scheme (Zhao and Carr, 1997, **ZC97**)

## Major features

1. **Cloud water and cloud ice** are prognostically calculated with one predictive variable( $q_c/q_i$ )
2. **Precipitation** diagnosed from cloud mixing ratio( $q_c/q_i$ ) and falling to next level immediately.

- 大氣中產生**雲滴**或是**雲冰**, 由溫度決定。
- 大氣中不會有**雨滴**或**雪**的停留。



## Wsm5 微物理方案 (Hong et al 2004)

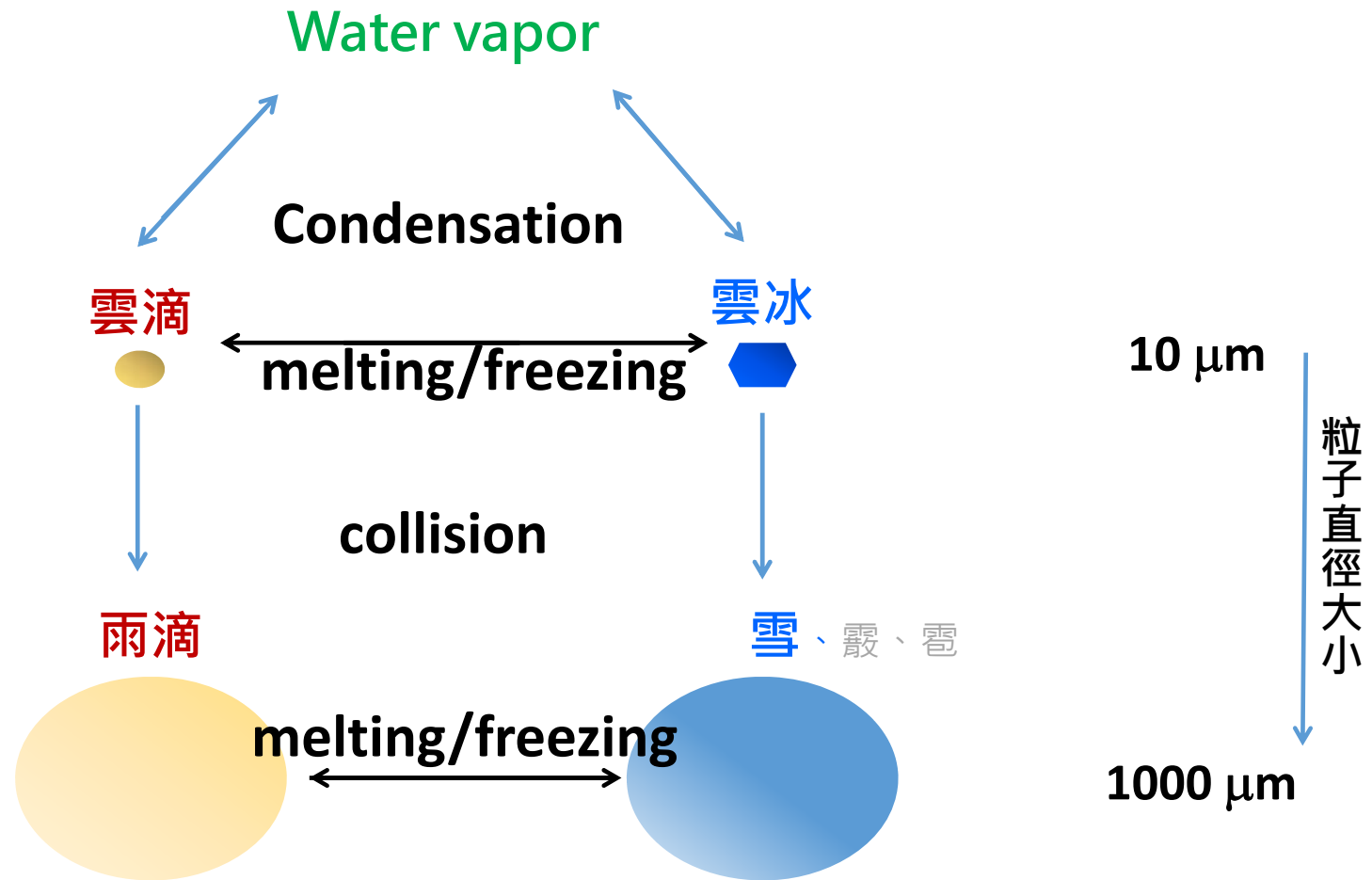
- 總體水物參數法(Bulk method)

假設水物粒子的粒徑-個數濃度分佈，為一連續函數

- 單矩量(single moment)

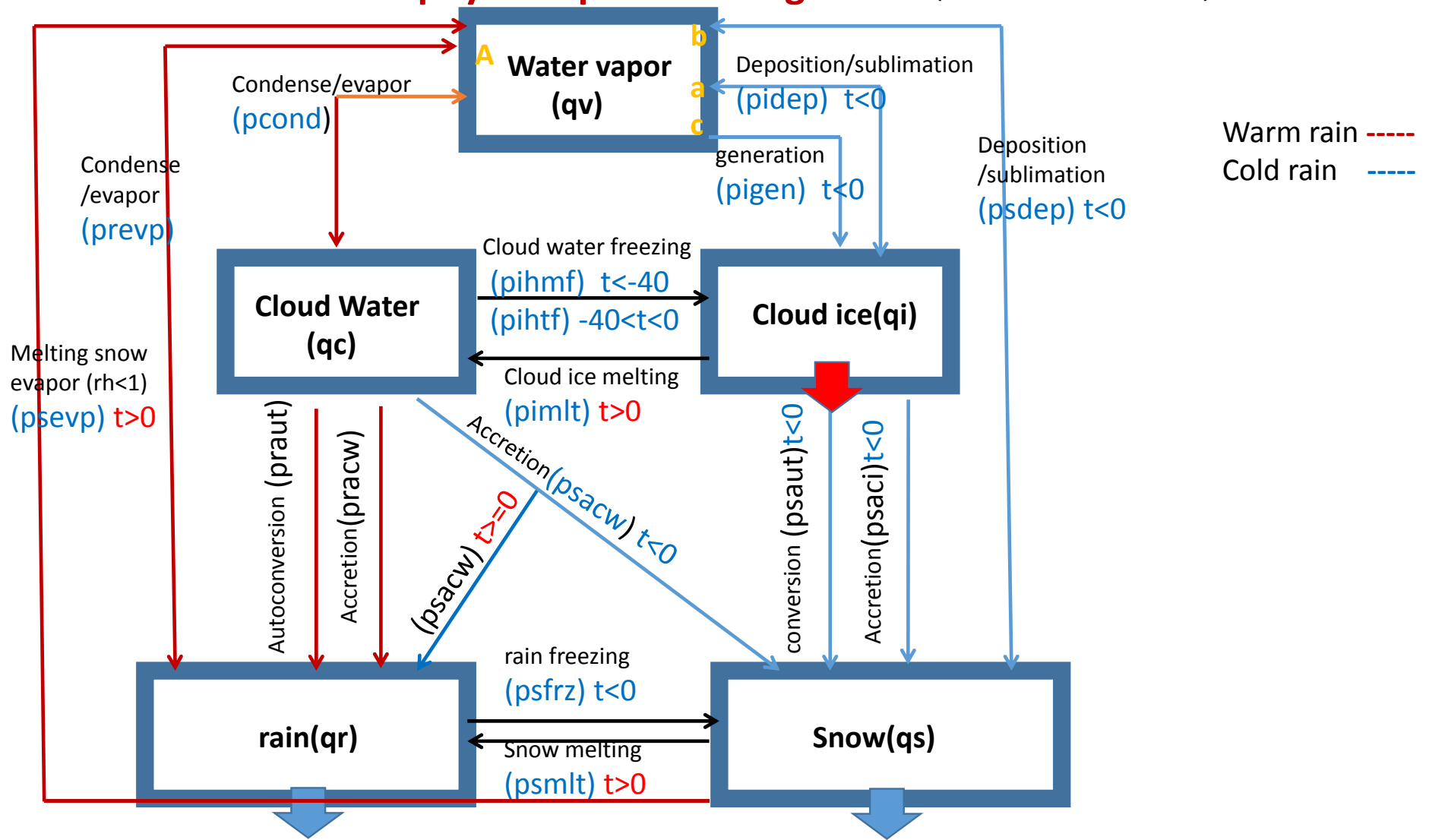
預報量--水物粒子的質量混和比

# WSM5 考慮的雲與降水粒子



# WSM5 microphysical process diagram

(ref to Lin et al 1983)



# CWBGFS 模式簡介

**Model : CWB-GFS T511L60**

💧 **Horizontal : ~ 25 km , vertical : 60 layers**

💧 **Physics :**

<b>Radiation</b>	<b>RRTMG scheme</b>
<b>Cumulus</b>	<b>New Simplified Arakawa-Schubert (Pan and Wu , 1994)</b>
<b>Large Scale Precipitation</b>	<b>Predict cloud water scheme (Zhao and Carr 1997)</b>
<b>Shallow Convection</b>	<b>Li and Young (1993)</b>
<b>PBL</b>	<b>First-order nonlocal scheme (Troen and Mahrt 1986)</b>
<b>Surface Flux</b>	<b>Similarity theory (Businger 1971)</b>
<b>Land Model</b>	<b>Noah Land model(4-layer)</b>
<b>Gravity Wave Drag</b>	<b>Palmer et al. (1986)</b>

- **Experiment design**

- ops : CWB GFS with ZC97 scheme
- **wsm5** : CWB GFS with **wsm5 scheme**

### Case study

- Initial 00 UTC 01 Jan, 2013

### Full cycle run

- 6-hr cyclic frequency
- 01-14 Jan, 2013
- Evaluation for 5-day prediction at 00/12 UTC run

- **Verification truth**

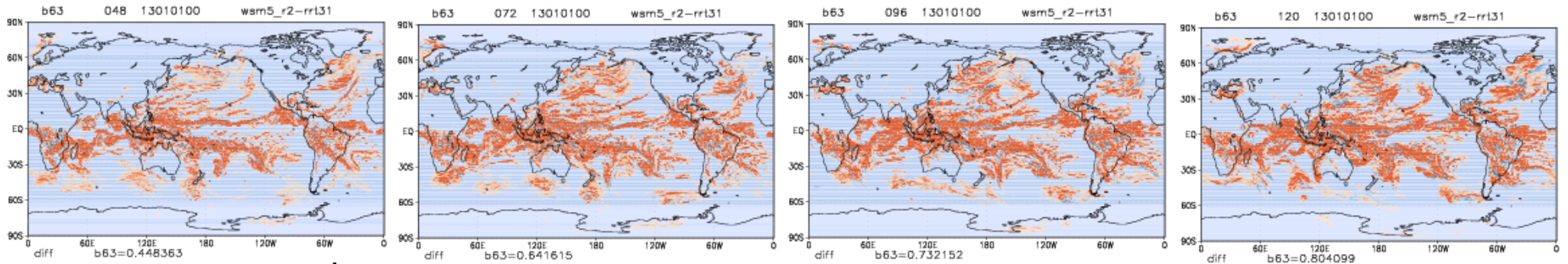
Era-Interim(0.75°x0.75°)



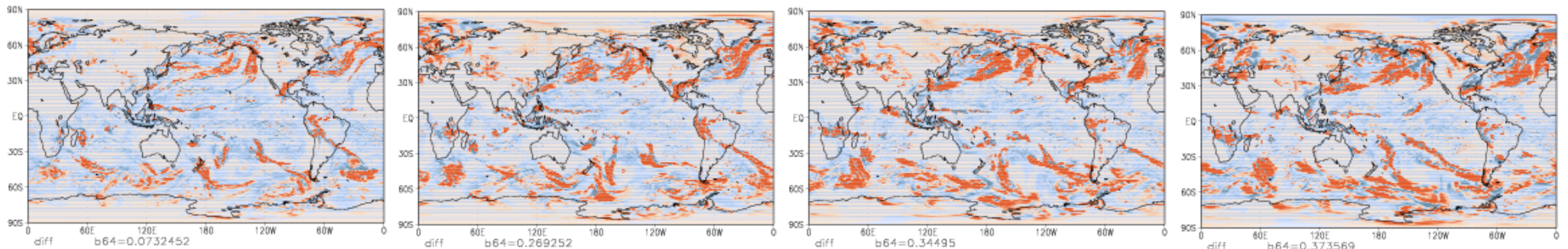
● **問題:** 降水增加，且差值隨預報時間變大

## Precipitation difference(wsm5-ops)

**Cup** + 0.448 mm/day      + 0.641 mm/day      + 0.732 mm/day      + 0.804 mm/day



**Lsp** + 0.073 mm/day      + 0.269 mm/day      + 0.344 mm/day      + 0.373 mm/day



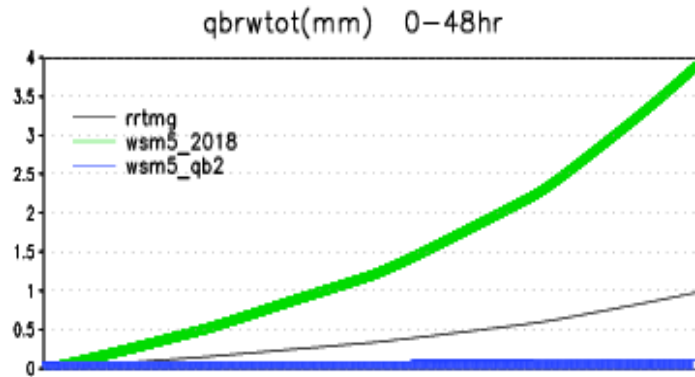
048 hr

072 hr

096 hr

120 hr

## Q borrow 累積量



Ops -----  
 Wsm5 -----  
 Wsm5\_qb -----

$$Q(\text{Ops}) : qv + qi / qc$$

$$Q(\text{Wsm5}) : qv + qi + qc + qs + qr$$

## 修改 Negative q 處理策略(qb)

- ops: 1. 向下層借  
 2. 最底層若有負值則無條件轉為正值

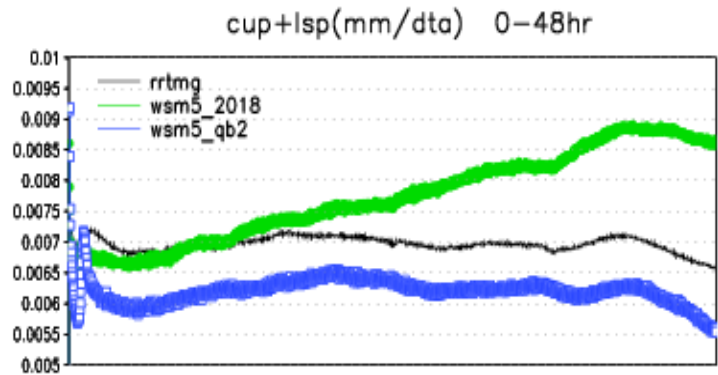
**總水量不保守!!**

- 修改: 1. 負值轉正，並將借量累計(qb)  
 2. 將正值點總和累加(qp)  
 3. 將正值點等比例降低，以維持總量保守。  
 $(q * (qp - qb) / qp)$

## Impact of Negative q adjust policy

- 降水率維持平穩
- 水物粒子混和比降低

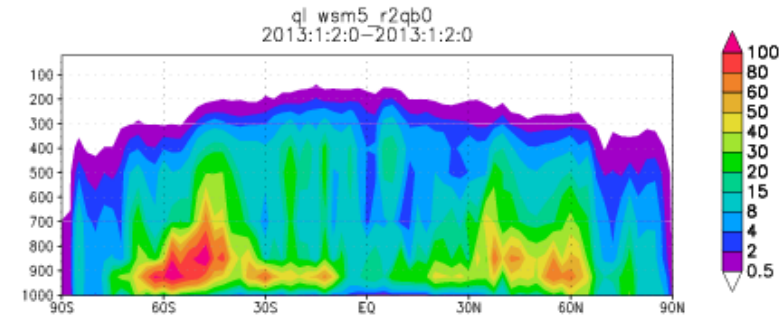
## Precip(mm/dta)



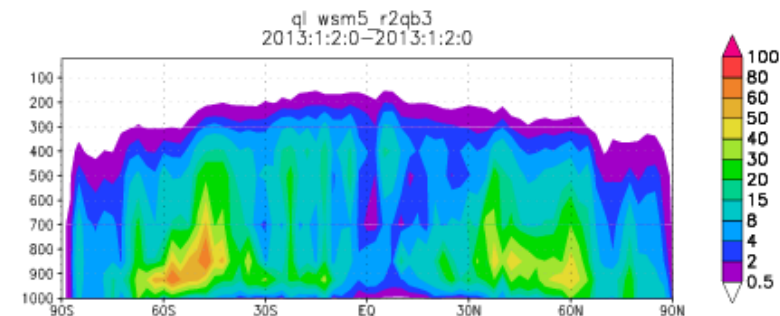
WSM5

WSM5\_qb

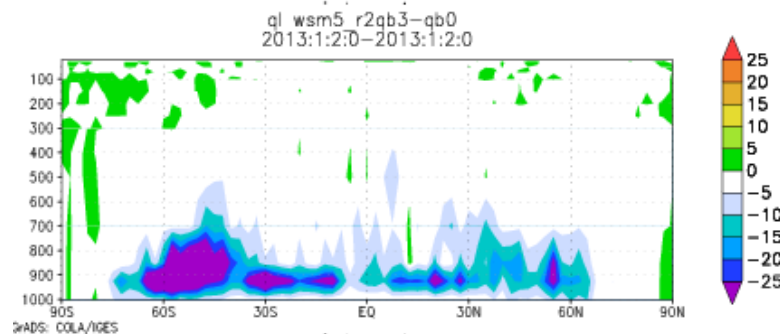
## Total ql (24h fcst) zonal mean , y-z profile



WSM5  
(qc+qi+qr+qs)



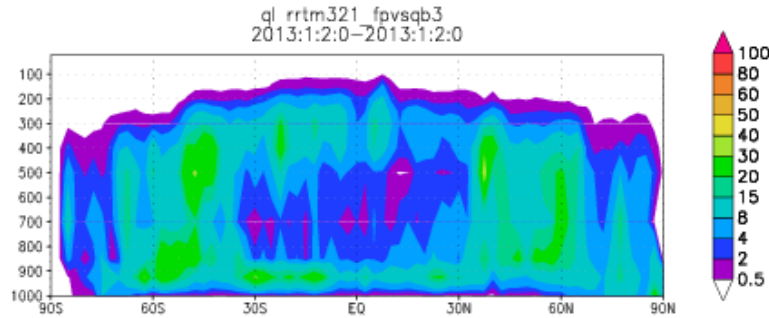
WSM5\_qb  
(qc+qi+qr+qs)



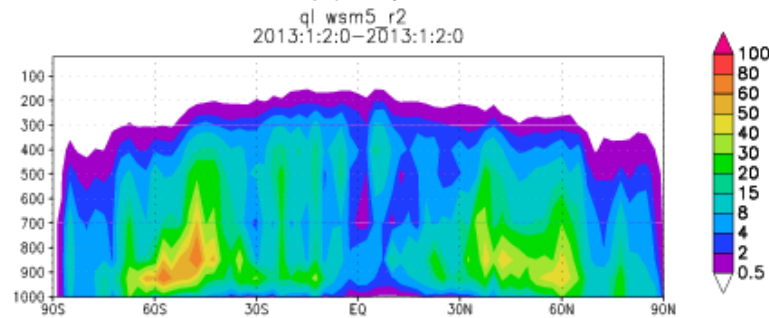
WSM5\_qb  
- wsm5

# Total ql (mg/kg) 24h fcst

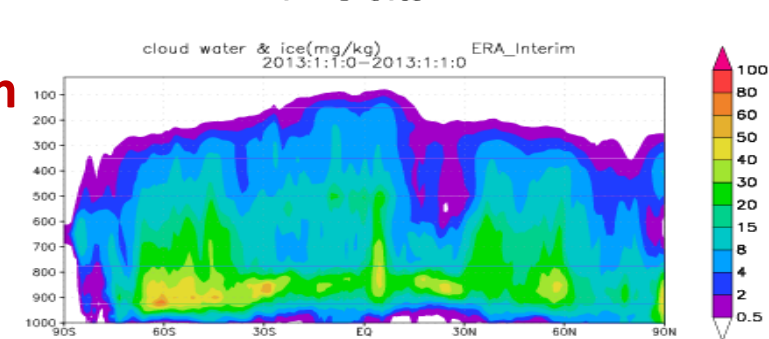
## zonal mean , y-z profile



ops  
(qc/qi)

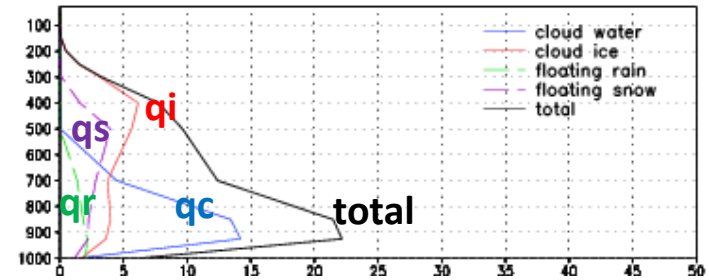
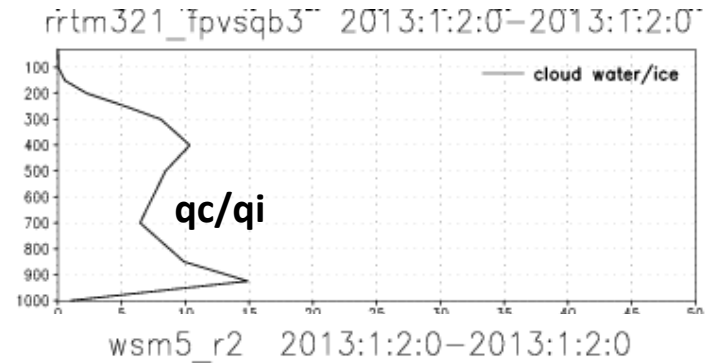


WSM5\_qb  
(qc+qi+qr+qs)

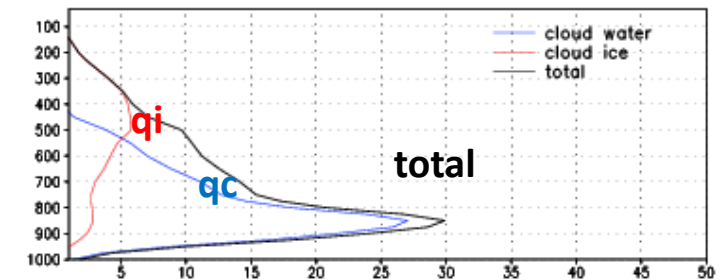


ERA\_interim  
(qc+qi)

## global mean , z profile

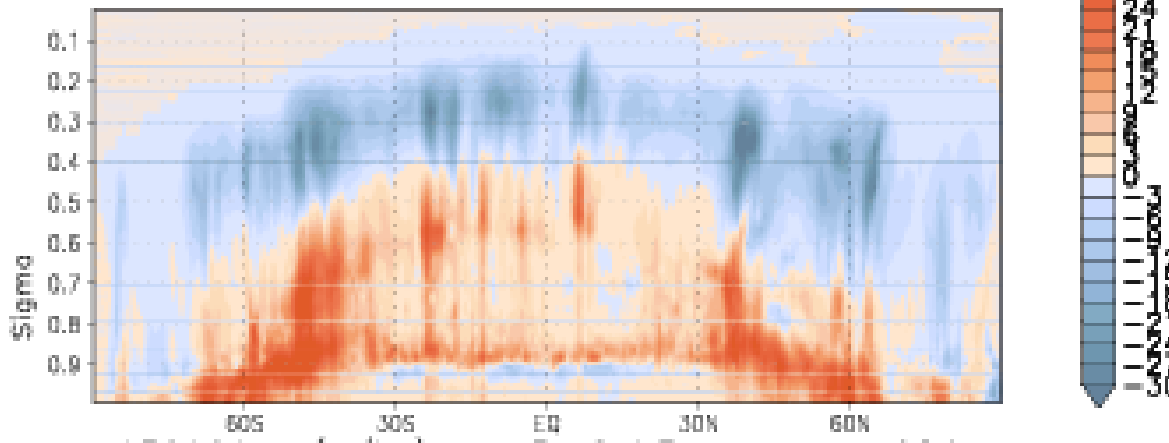


ERA-interim 2013:1:2:0-2013:1:2:0



## Total ql(mg/kg) Wsm5 - ops

130101 qr(mg/kg) wsm5\_r2qb3-ops 024



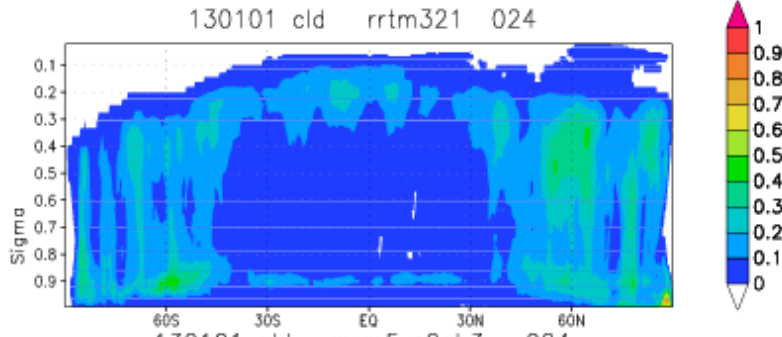
## Wsm5 impact on ql

$$ql = qc + qi + qs + qr$$

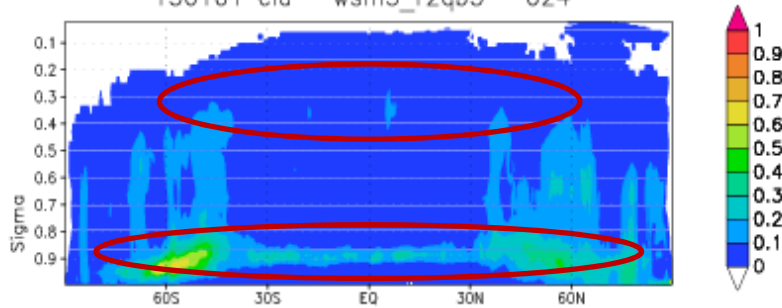
- ql 高層減少，中低層增加

ops

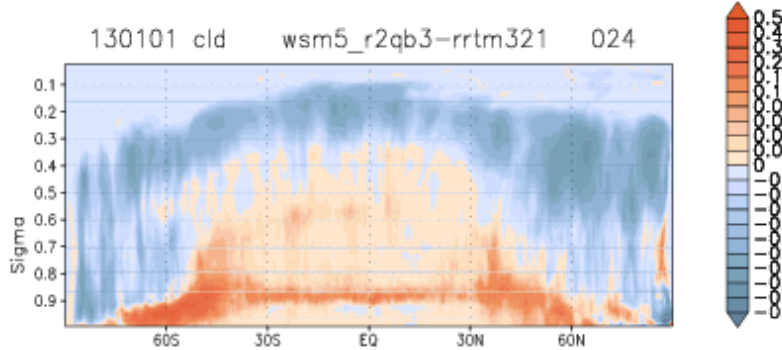
## Cloud fraction



Wsm5\_qb



Wsm5\_qb-ops



## Wsm5 impact on cloud

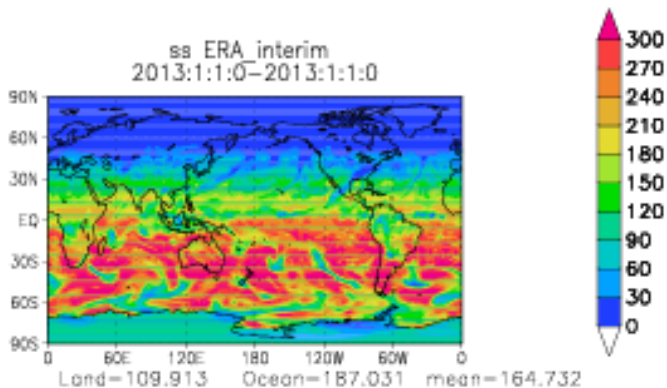
- 高層雲量減少
- 低層雲量增加

## Cloud fraction

$$c = rh^{1/4} \left( 1 - \exp \left( \frac{-100ql}{((1 - rh)qs)^{0.49}} \right) \right)$$

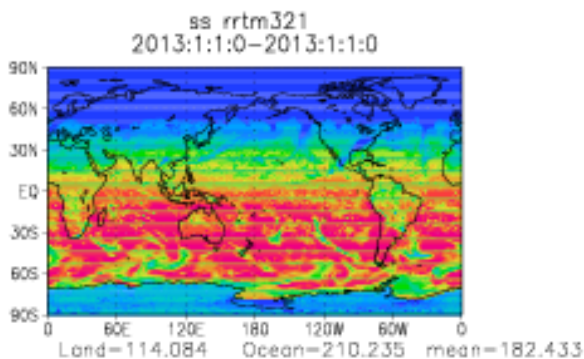
$ql = qi + qc + rs + qr$  (Xu and Randall 1996)

Era\_interim  
164.7 W/m<sup>2</sup>

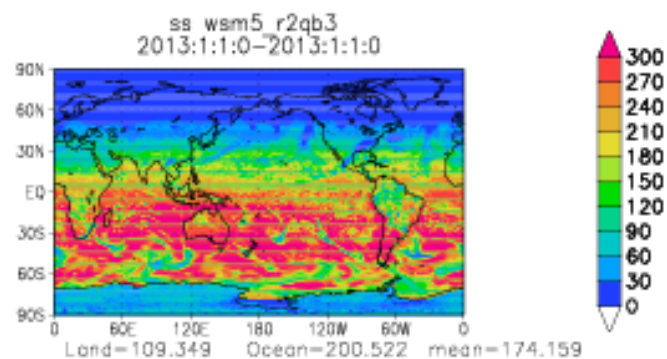


## Surface net solar rad (W/m<sup>2</sup>) 2013/1/1 day 1 fcst

ops  
182.4 W/m<sup>2</sup>

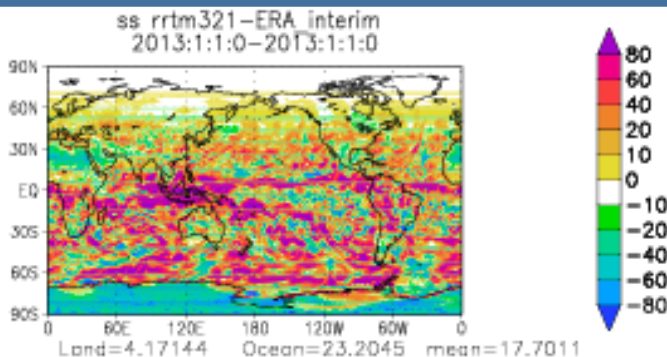


Wsm5\_qb  
174.1 W/m<sup>2</sup>



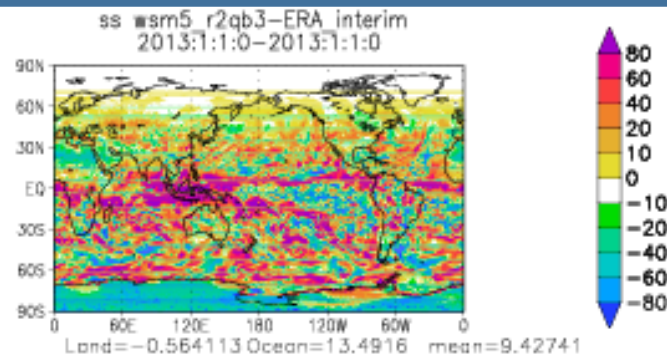
ops-Era  
+17.7 W/m<sup>2</sup>

>> SS過大

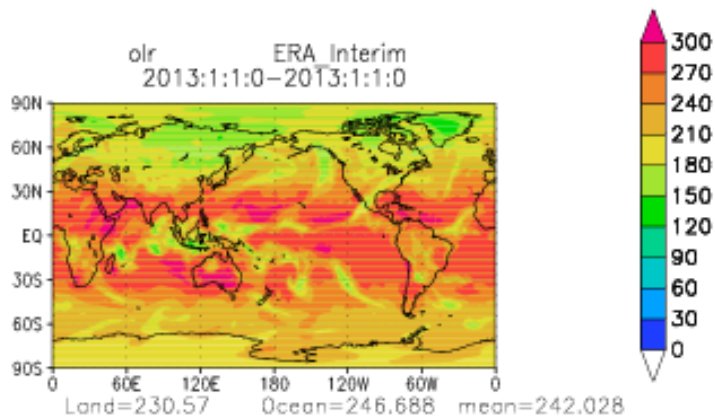


Wsm5\_qb-Era  
+9.4 W/m<sup>2</sup>  
伴隨低層雲量增加,  
SS 顯著減少

>> 較趨近ERA

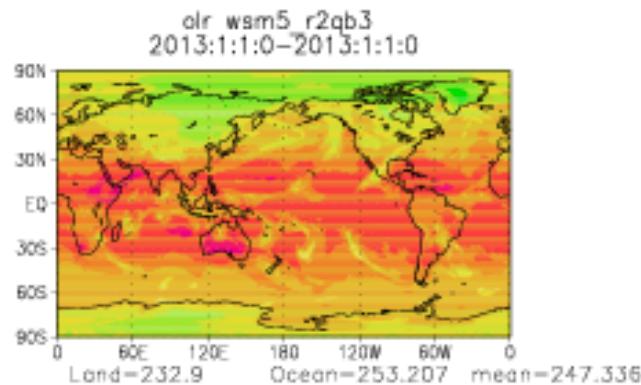
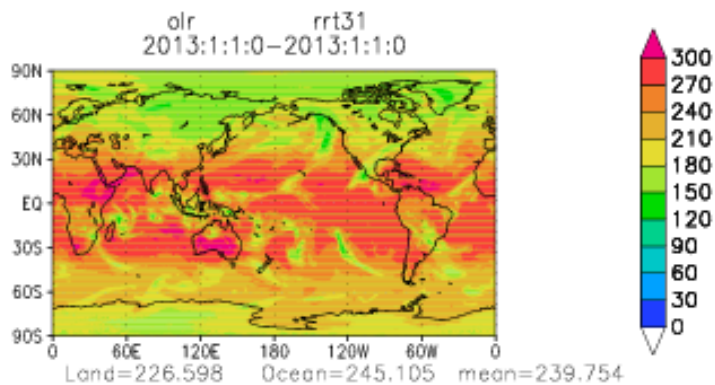


Era\_interim  
242.0 W/m<sup>2</sup>



## Outgoing Longwave Rad(W/m<sup>2</sup>) 2013/1/1 day 1 fcst

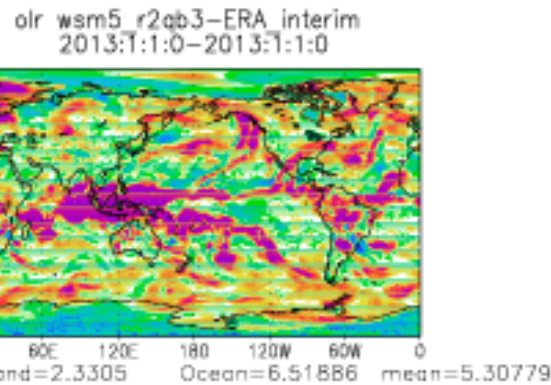
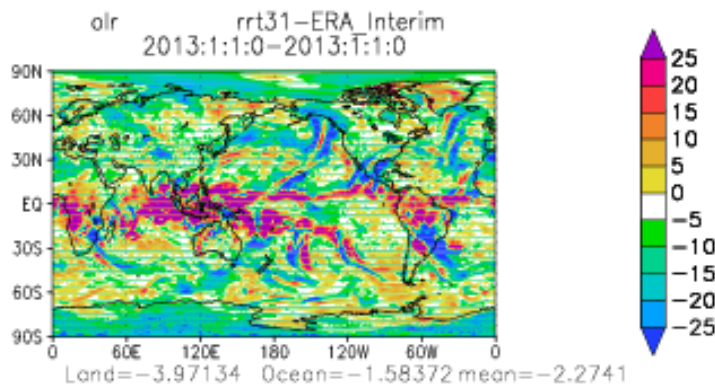
ops  
239.7 W/m<sup>2</sup>



Wsm5\_qb  
247.3 W/m<sup>2</sup>

ops-Era  
-2.27 W/m<sup>2</sup>

>> OLR 偏少



Wsm5\_qb-Era  
+5.3 W/m<sup>2</sup>  
伴隨高層雲量減少  
Olr大幅增加

>>有增加過多趨勢

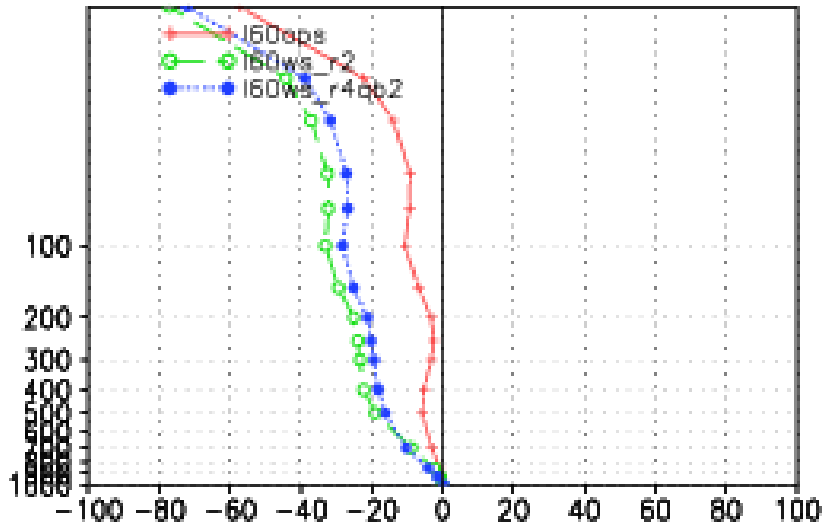


ops -----  
 wsm5 -----  
 wsm5\_qb -----

T511L60  
 2013/01/01-14  
 day 5 FCST NA(20N-80N)

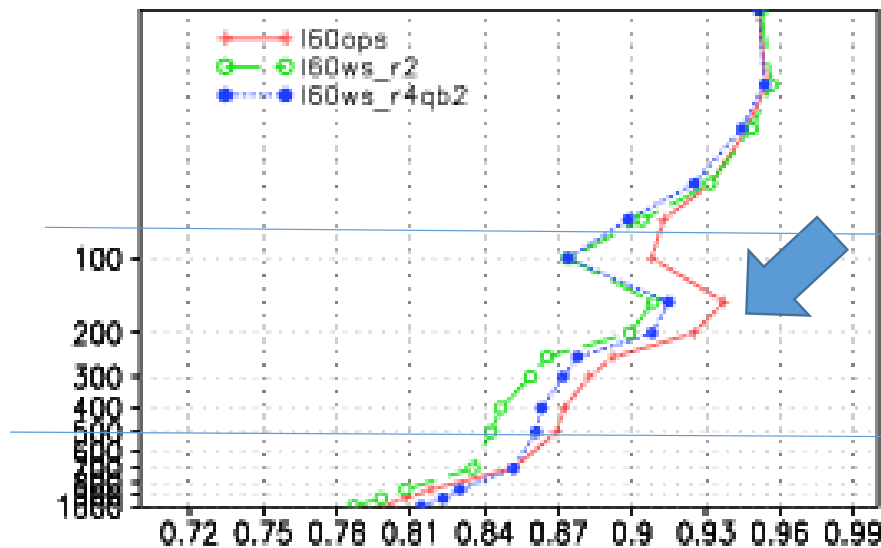
### FMH 負偏差增加

2013 Jan fmh 5 day fcst - NA  
 2013:1:1:0-2013:1:14:0



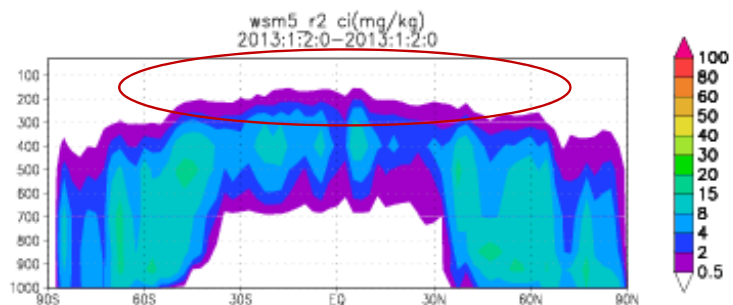
### ACH 顯著降低

2013 Jan ACH 5 day fcst - NA  
 2013:1:1:0-2013:1:14:0

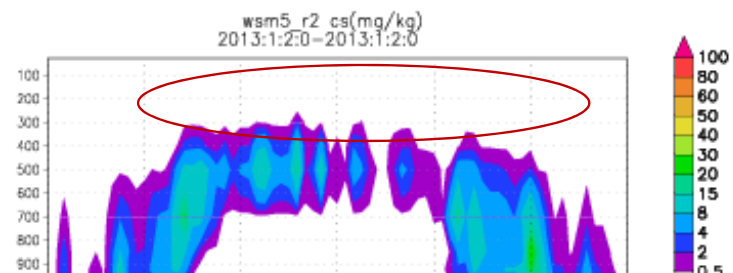


# WSM5\_qb 130101 24hr FCST

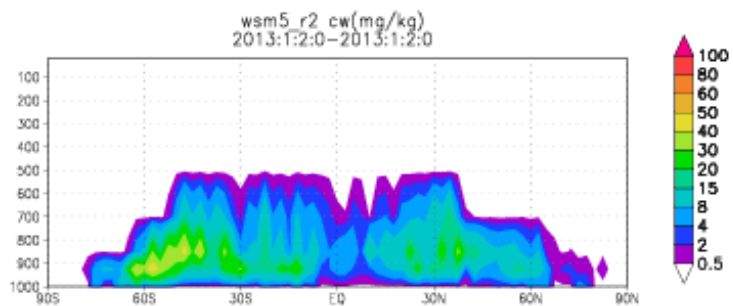
Qi(mg/kg)  
Cloud ice



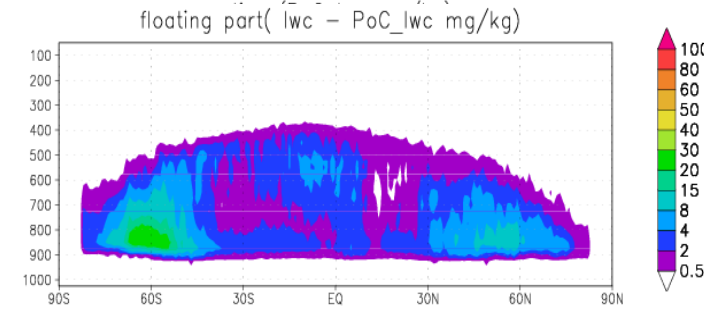
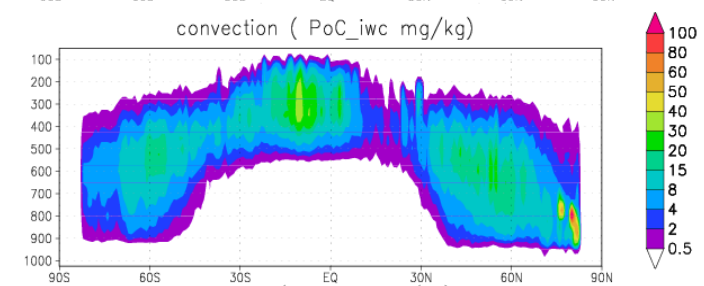
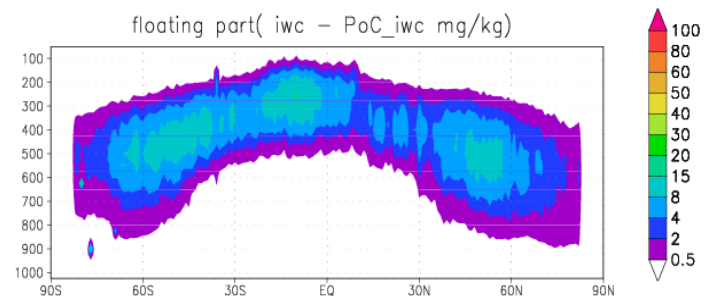
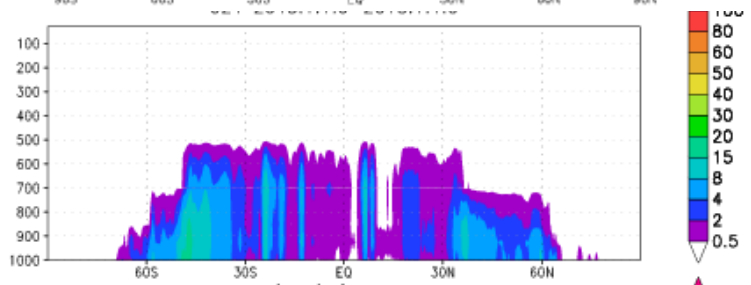
Qs(mg/kg)  
snow



Qc(mg/kg)  
Cloud water



Qr(mg/kg)  
rain



**CLDSAT DATA(Jan. 2013)**

➤ 大氣中高層水物粒子偏少

## 調整測試:

- 高雲過少

- 減小雲冰的終端速度(增加雲冰於高層的停留)

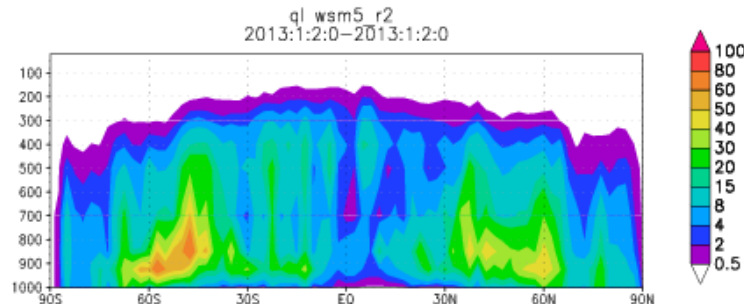
Vice = 0.01 m/s ( wsm5\_qbvi01)

$$V_t(\text{ice}) = 3.29(\rho q_i)^{0.16}$$

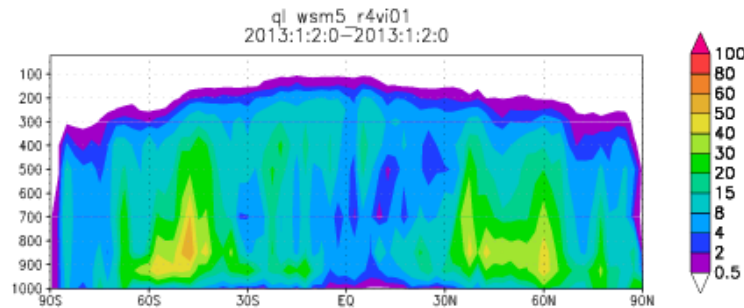
Original Wsm5 雲冰終端速度

# Wsm5\_qbvi01 impact On ql (mg/kg)

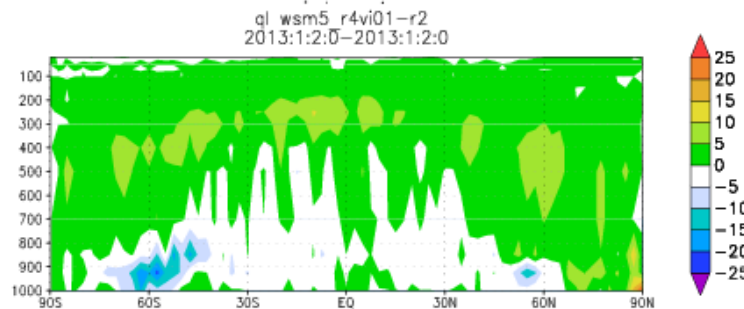
**Total ql zonal mean , y-z profile**



Wsm5\_qb  
qi+qc+qs+qr

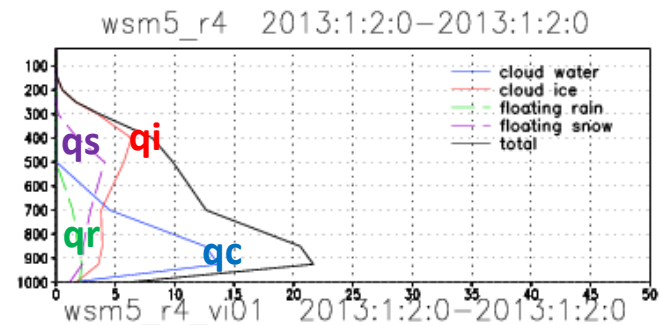


Wsm5\_qbvi01  
qi+qc+qs+qr

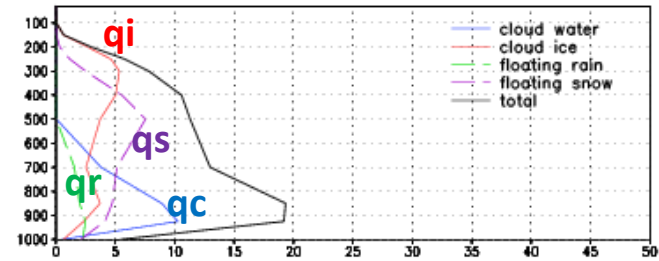


Diff  
Wsm5\_qbvi01  
-wsm5\_qb

**Total ql global mean , z profile**



Wsm5\_qb

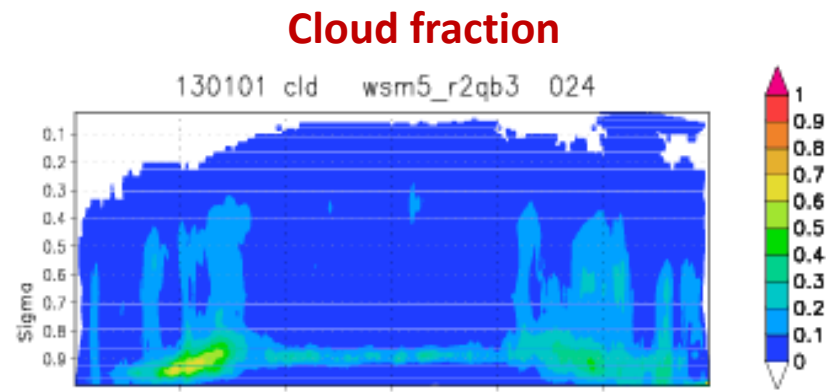


Wsm5\_qbvi01

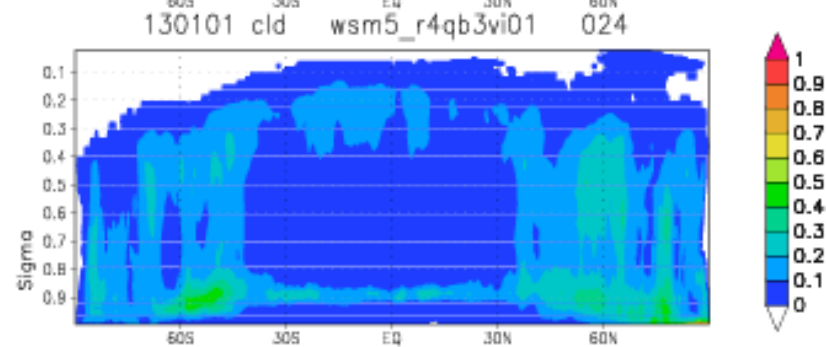
- qi (cloud ice) 分佈往更高層延伸
- qs (snow) 量值增加

# Wsm5\_qbvi01 impact On cloud fraction

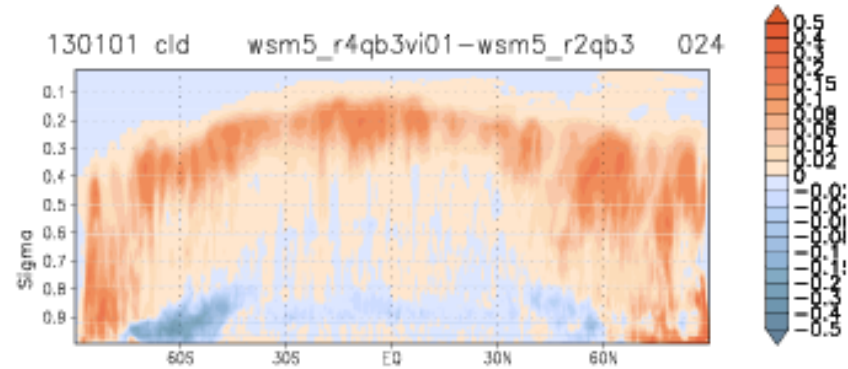
Wsm5\_qb



Wsm5\_qbvi01



Diff  
Wsm5\_qbvi01  
-wsm5\_qb

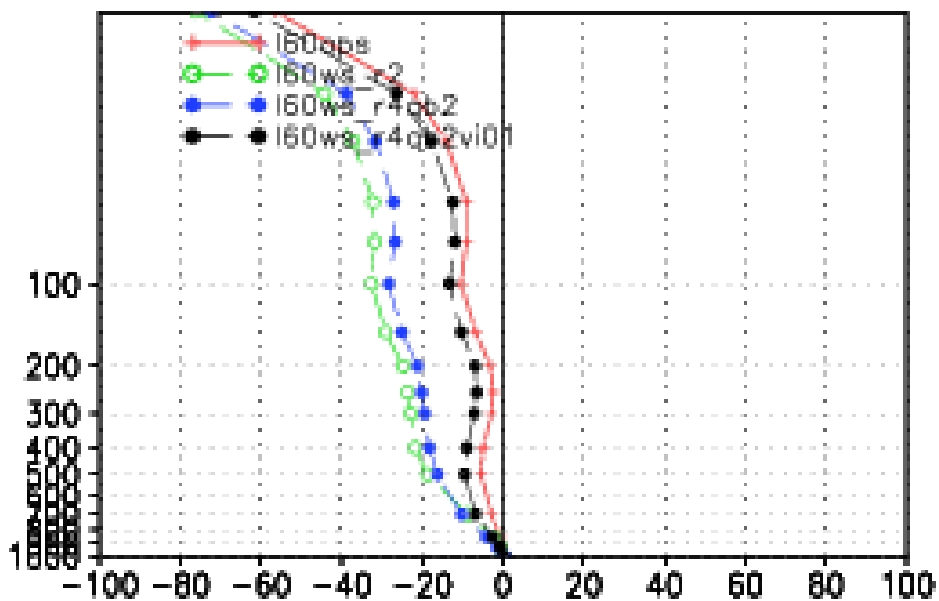


**ops**                    - - - - -  
**wsm5**                    - - - - -  
**wsm5\_qb**                - - - - -  
**wsm5\_qbvi01**         - - - - -

T511L60  
 2013/01/01-14  
 day 5 FCST NA(20N-80N)

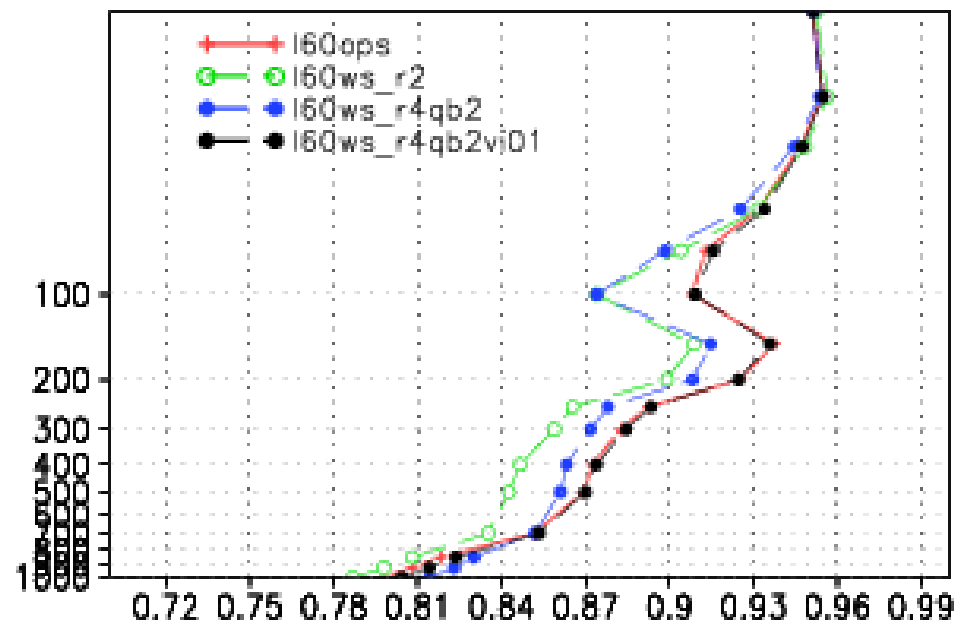
**FMH**

2013 Jan fmh 5 day fcst - NA



**ACH**

2013 Jan ach 5 day fcst - NA



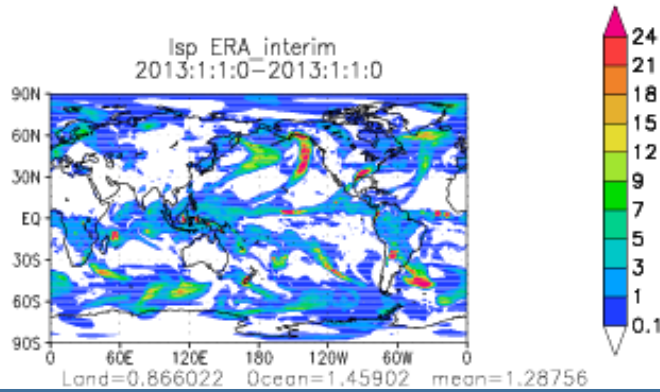
## 待解決問題:

- 網格尺度降水偏少
  - wsm5設定的凝結條件，  
是要求格點平均濕度達飽和。

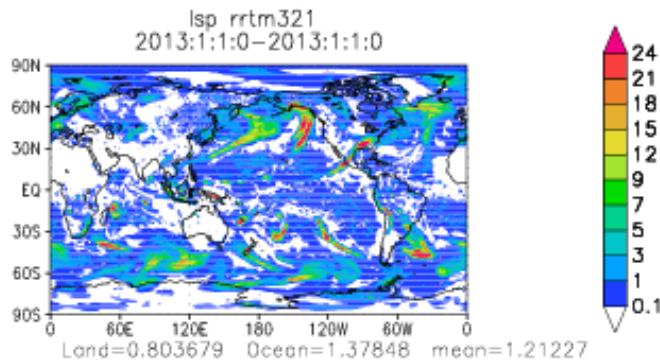
此條件對較粗網格(25km)模式是否適合?

# Grid-scale precipitation (mm/day) 2013/1/1 day 1 fcst

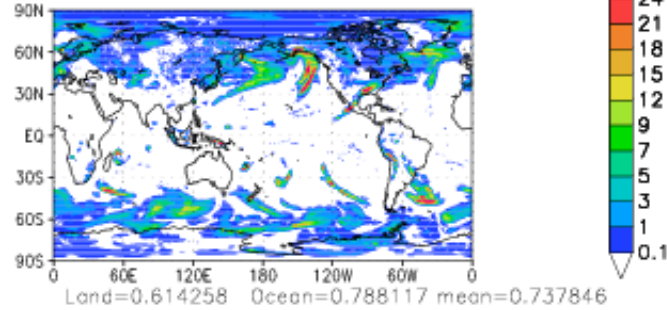
Era\_interim  
1.287 mm/day



ops  
1.212 mm/day

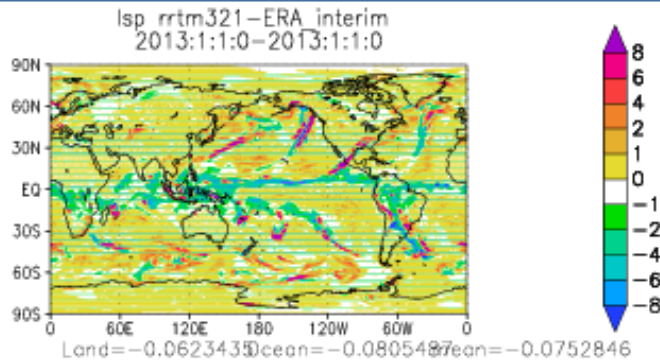


lsp wsm5\_r2qb3  
2013:1:1:0-2013:1:1:0

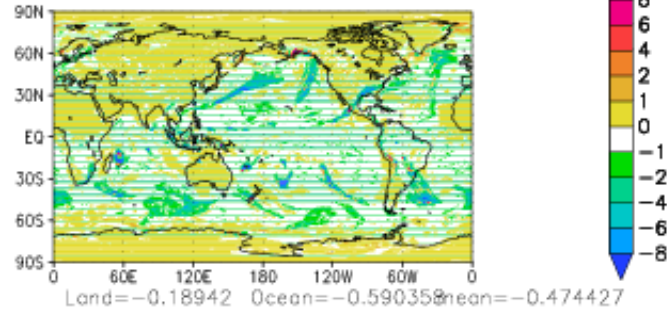


Cwb\_wsm5  
0.737 mm/day

Ops-era  
-0.07 mm/day



lsp wsm5r2qb3-rrtm321  
2013:1:1:0-2013:1:1:0



Wsm5\_qb-ops  
-0.47 mm/day

網格降水偏少

洋面網格降水減少  
偏少誤差增加



## WSM5 評估與建議

- 水物粒子分布(雲)有較為合理的基本垂直結構

多項微物理過程可提供較大調整彈性，

得到最佳輻射通量 ( olr,ss )

**但是**

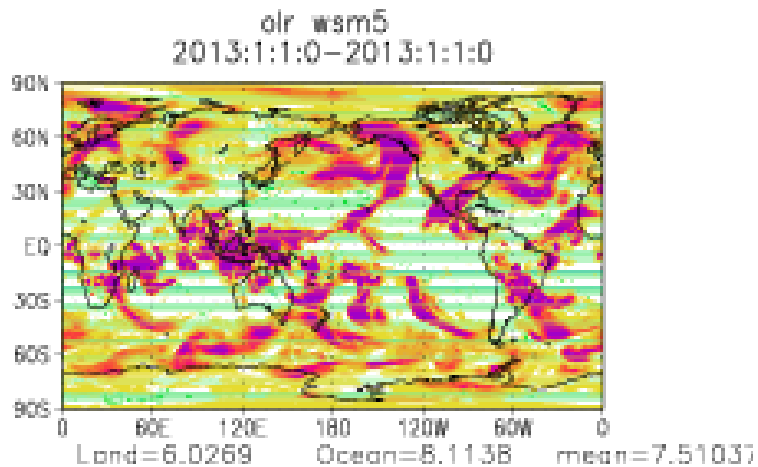
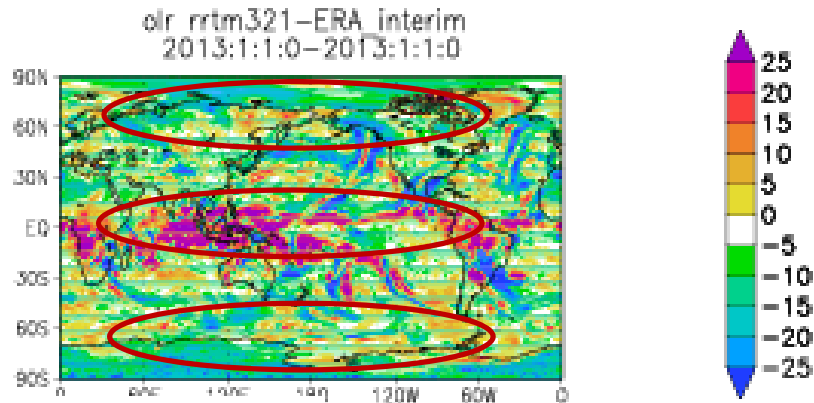
- 模式必需水量保守
- 解析度必須提高，解決網格尺度降水效率偏低問題
- 時間耗費為原3/2，須有足夠電腦資源

*The end -*

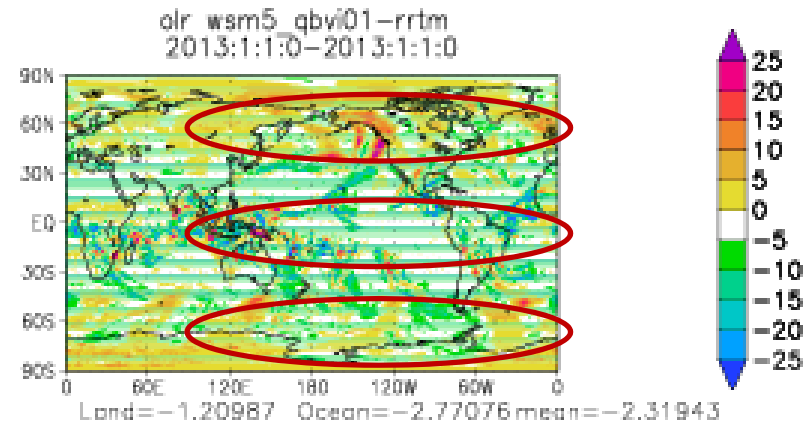
# Wsm5\_qbvi01 impact on

# Outgoing Longwave Rad(W/m2) 2013/1/1 day 1 fcst

Ops-Era

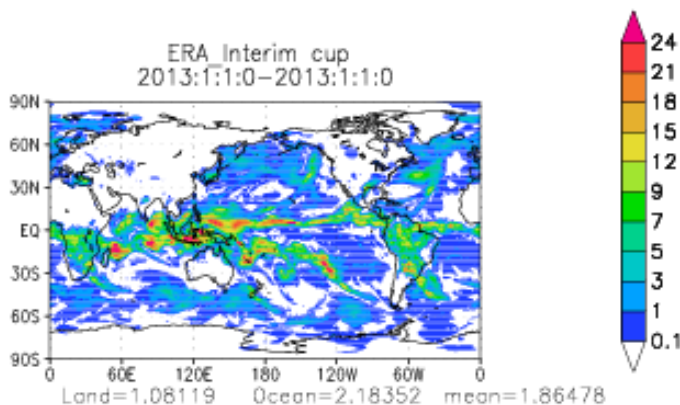


Wsm5-Ops



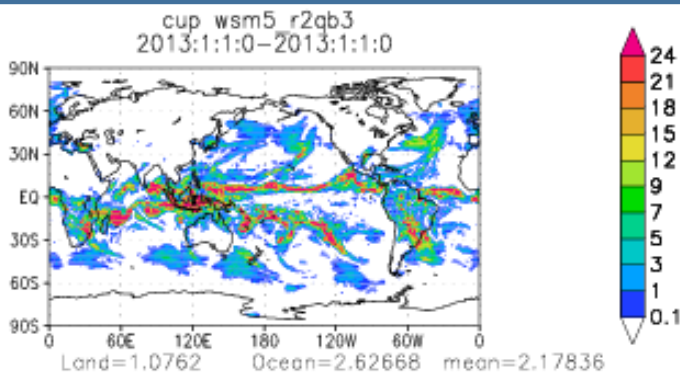
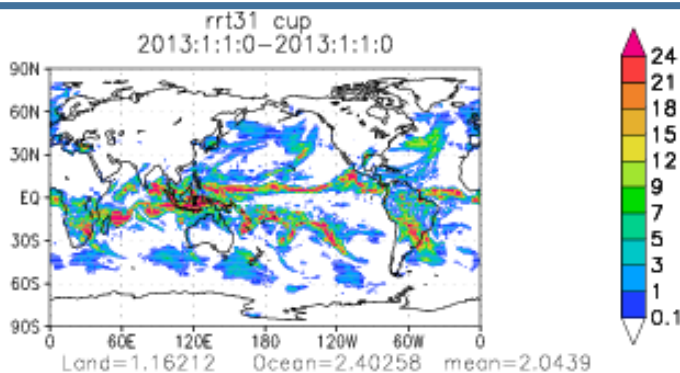
Wsm5\_viol-Era

Era\_interim  
1.864 mm/day



## Cumulus precipitation (mm/day) 2013/1/1 day 1 fcst

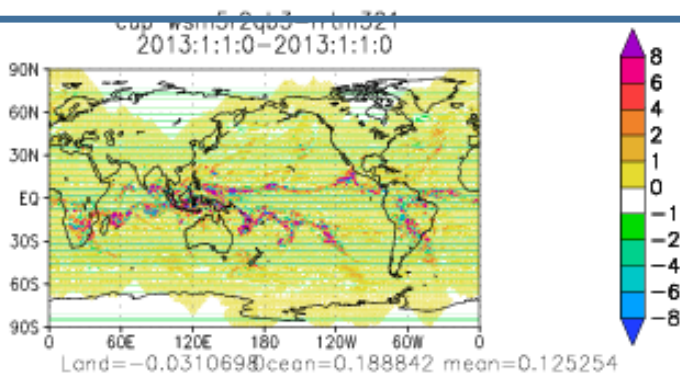
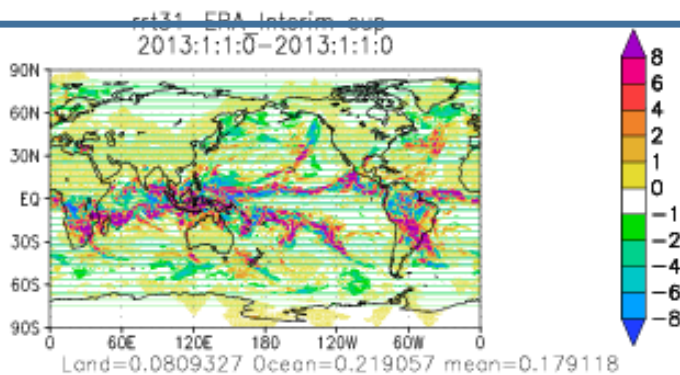
ops  
2.043 mm/day



Wsm5\_qb  
2.178 mm/day

Ops-era  
0.17 mm/day

對流降水偏多



Wsm5\_qb-ops  
0.13 mm/day

對流降水普遍增加  
偏多誤差增加

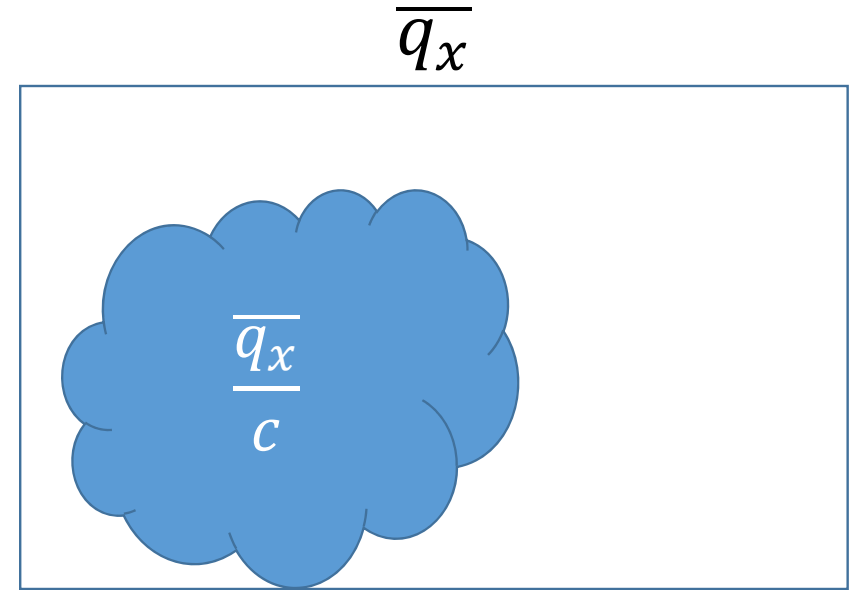
方法 1

$$\frac{d\overline{q_x}}{dt} = f\left(\frac{\overline{q_x}}{c}\right) \times c$$

$\overline{q_x}$  : grid mean mixing ratio

$C$  : cloud fraction

$\frac{\overline{q_x}}{c}$  : in cloud mixing ratio



方法 2

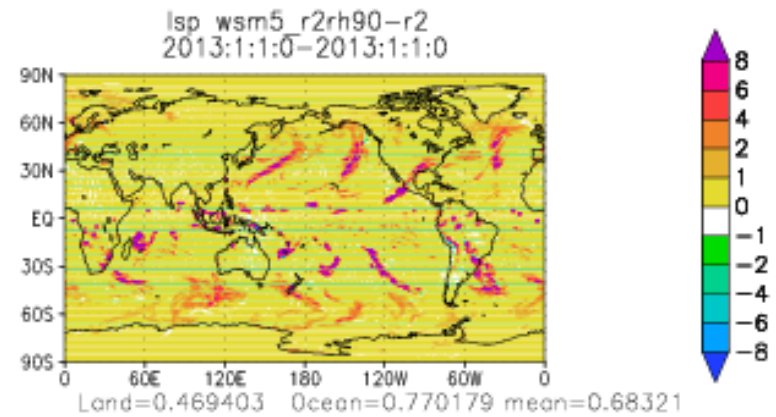
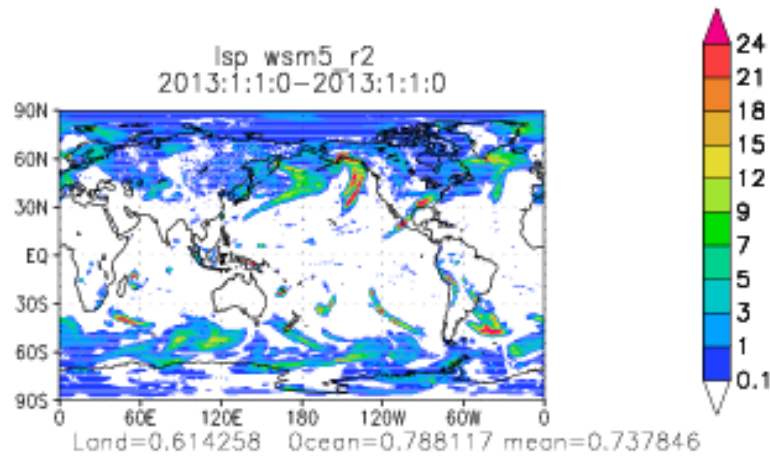
$$\overline{q_{s\text{modify}}} = \text{rh\_cri} \times \overline{q_s}$$

$\overline{q_s}$  : saturated mixing ratio

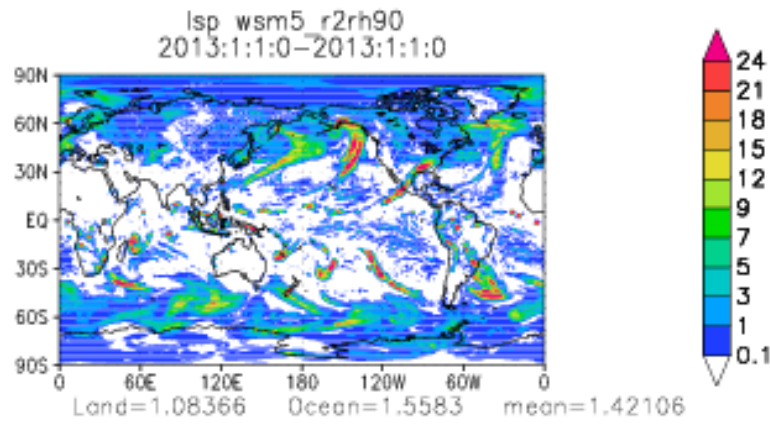
$$0 < \text{rhcri} < 1$$

# Rh=0.9 impact on Isp

Wsm5  
0.74 mm/day



Wsm5\_rh90  
1.42 mm/day

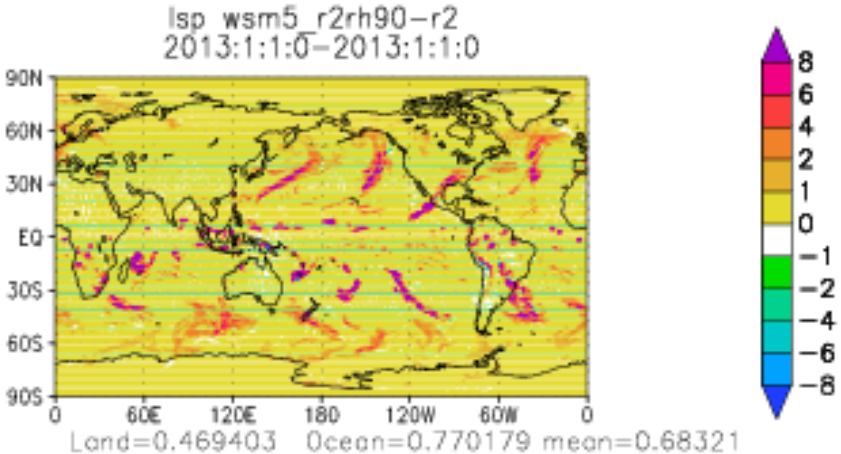


Diff = 0.68 mm/day  
rh90-wsm5

**Grid scale precip is increasing!!**

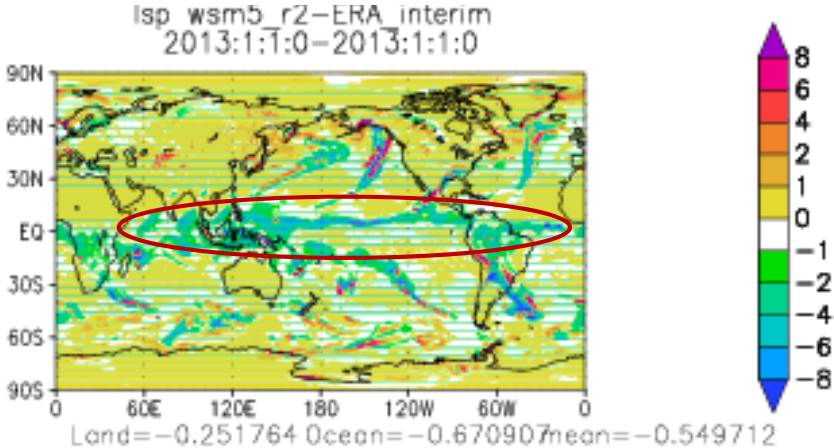
# Rh=0.9 impact on lsp

Wsm5\_rh90-wsm5



Lsp 全面增加

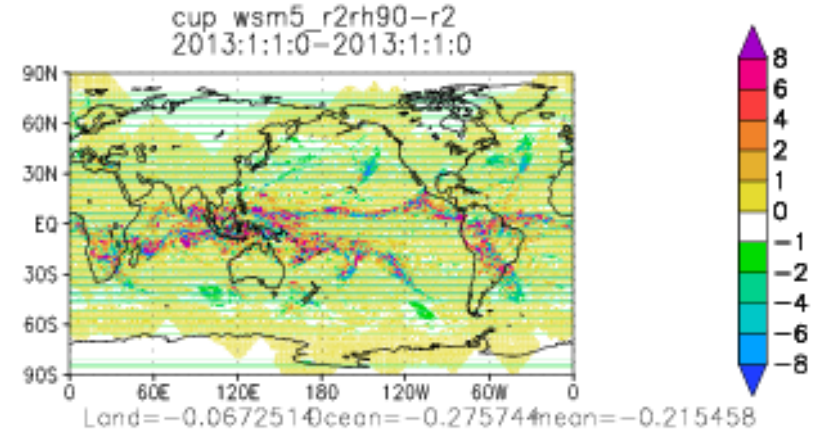
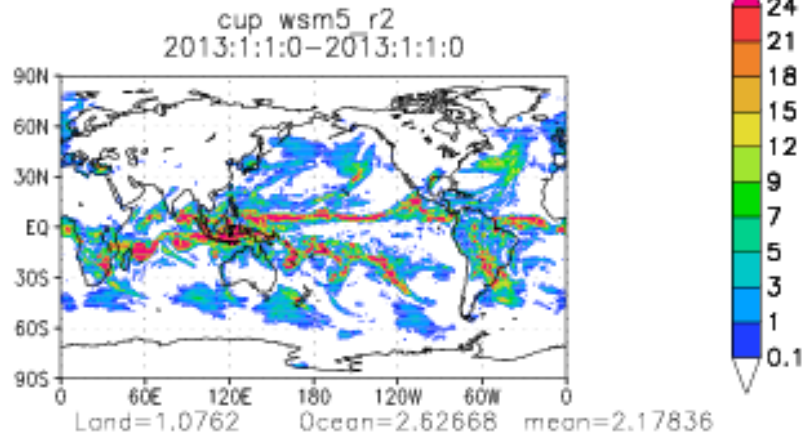
Wsm5-ERA\_interim



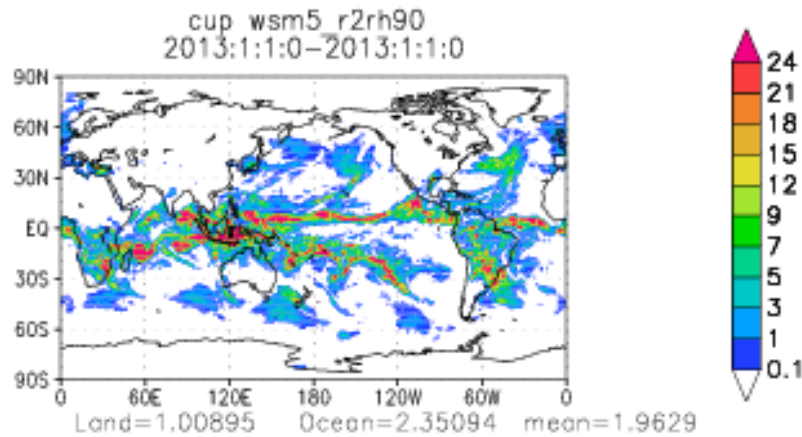
Lsp 主要偏少區域在熱帶

# Rh=0.9 impact on **cup**

Wsm5  
2.17 mm/day



Wsm5\_rh90  
1.96 mm/day



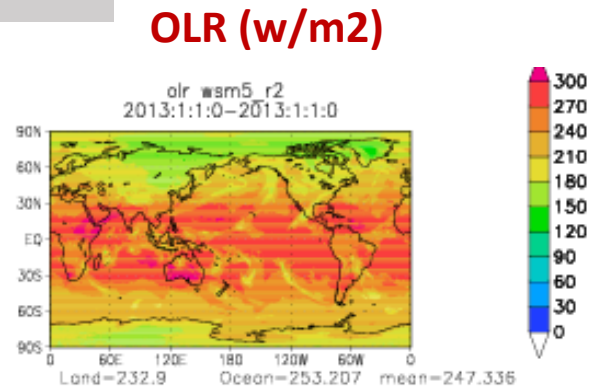
Diff = -0.215 mm/day  
rh90-wsm5

**Convective precip is decreasing!!**

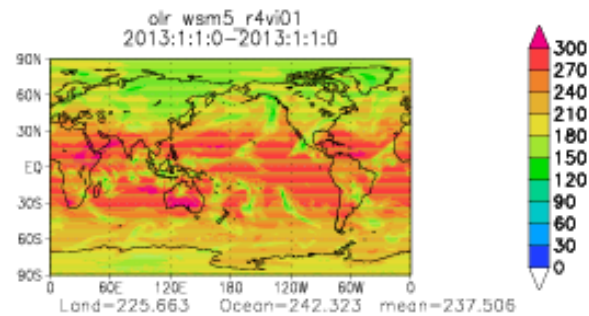


# Wsm5\_r4vi01 impact On OLR & cfr

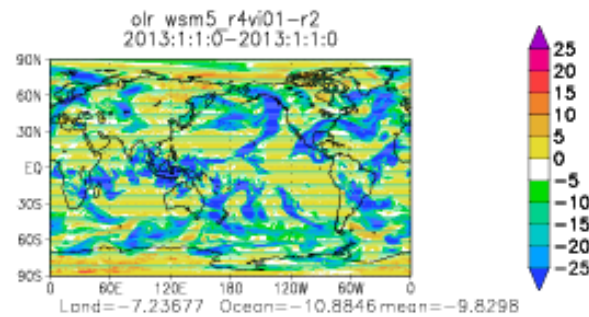
Wsm5\_r2  
247.3 W/m<sup>2</sup>



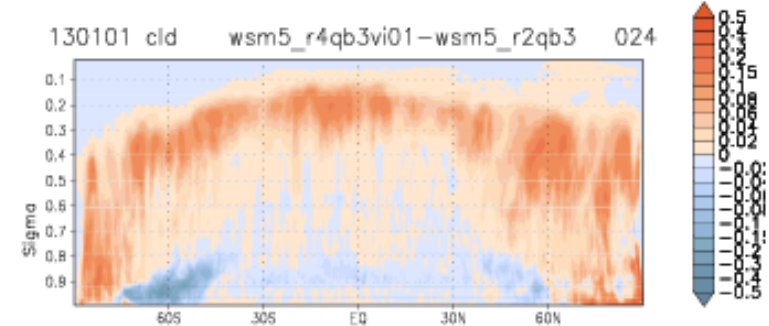
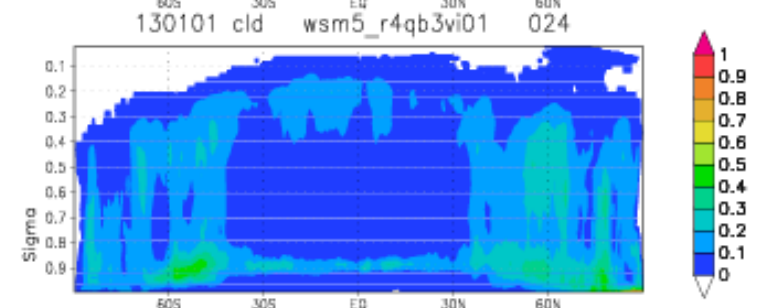
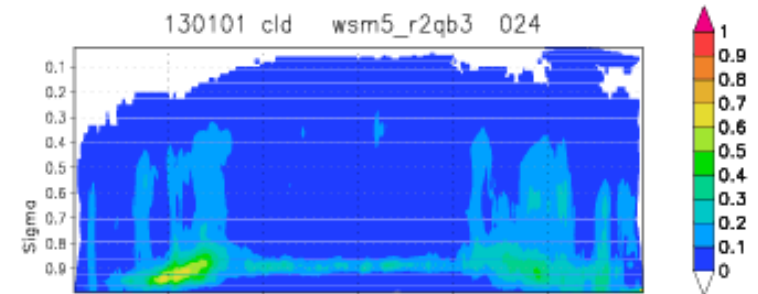
Wsm5\_r4vi01  
237.5 W/m<sup>2</sup>



Wsm5\_r4vi01  
-r2



## Cloud fraction



## Goddard bulk microphysical scheme (from WRF)

- Bulk-method, Single moment scheme
- 考慮的水物粒子  $q_c, q_r, q_i, q_s, q_g, q_h$  (2ICE, cloud ice & snow)

粒徑分布函數亦是M-P分布，但參數設定不同

沒有考慮cloud ice的沉降

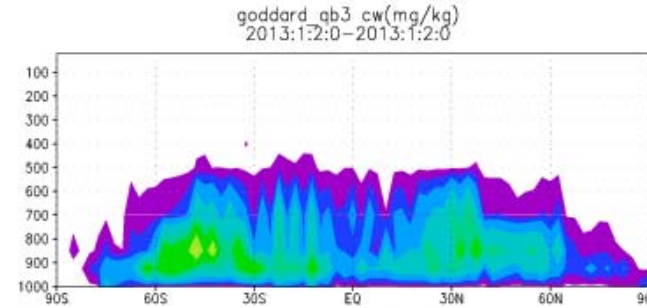
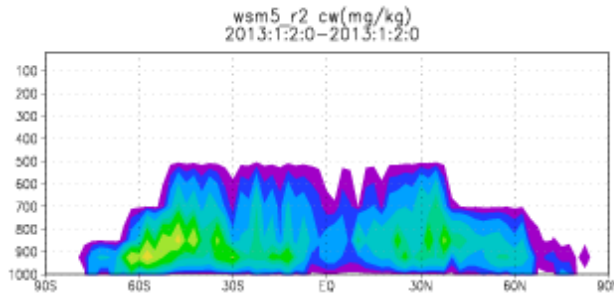
20130101 024

Yz profile

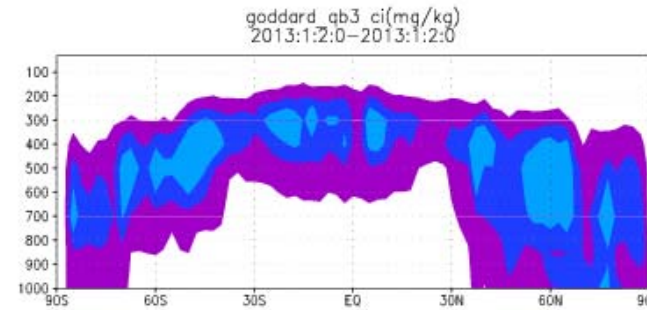
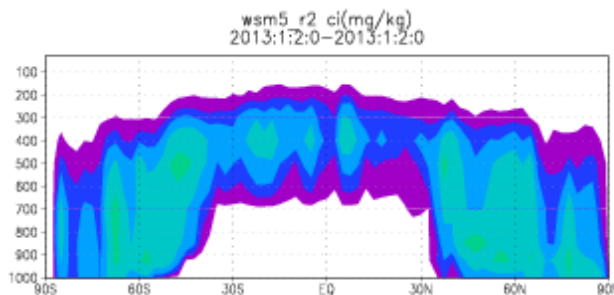
Wsm5

Goddard

qc  
Cloud water

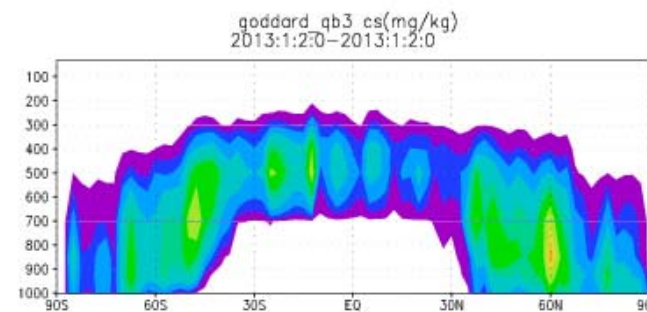
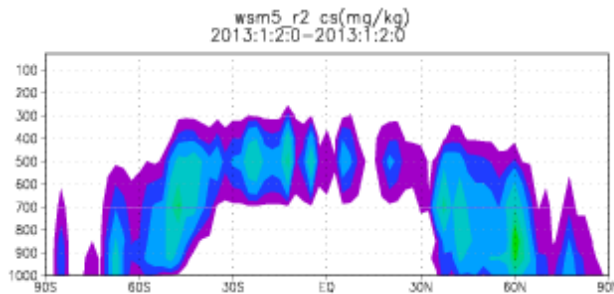


qi  
Cloud ice



Cloud Ice 較少

qs  
snow

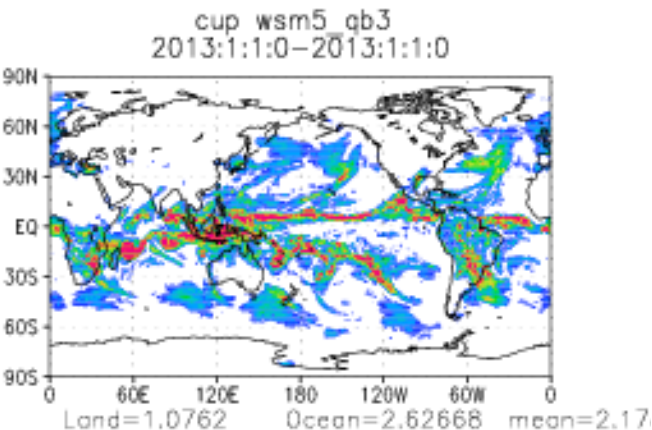


Snow 較多

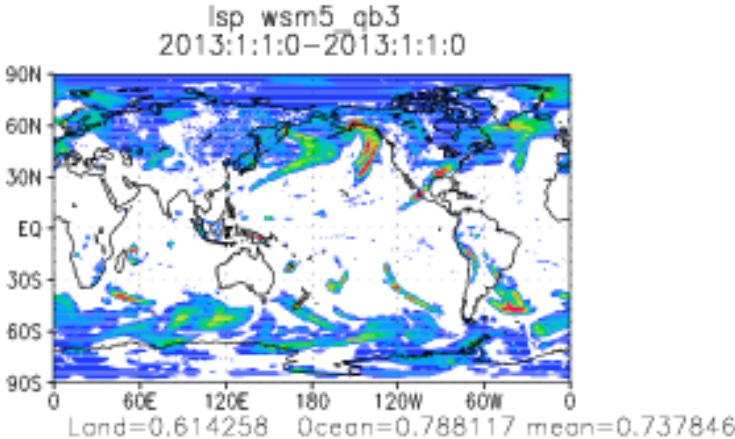
# Goddard scheme impact

Wsm5  
2.17 mm/day

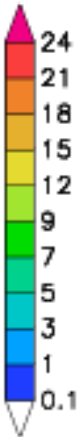
cup



lsp

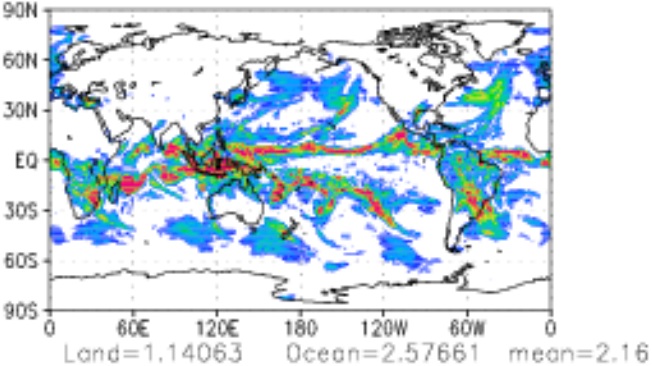


Wsm5  
0.73 mm/day

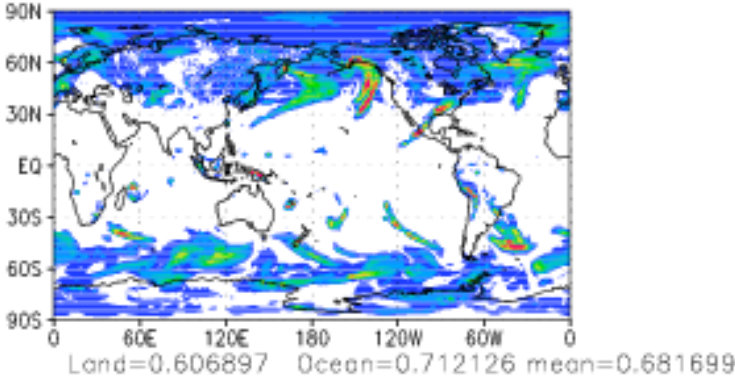


Goddard  
2.16 mm/day

cup goddard qb3  
2013:1:1:0-2013:1:1:0



lsp goddard qb3  
2013:1:1:0-2013:1:1:0



Goddard  
0.68 mm/day



Wsm5 (ref to Lin et al. 1983 ,Rutledge and Hobbs 1983)

**Modification** (Hong et al 2004)

**高層雲冰過多問題**  $N_{io} = 10^{-2} \exp(0.6(T_0 - T))$  (Fletcher 1962)

■ 冰核(nuclei)粒子濃度與冰晶(crystal)粒子濃度分開估算

- 假設冰核濃度為溫度的函數 *ice nuclei*  $N_{io} = 10^3 \exp(0.1(T_0 - T))$
- 假設雲冰濃度是雲冰質量混合比的函數

■ 加入雲冰的沉降作用  $V_t(\text{ice}) = 3.29(\rho q_i)^{0.16}$

## Cloud-radiation effect

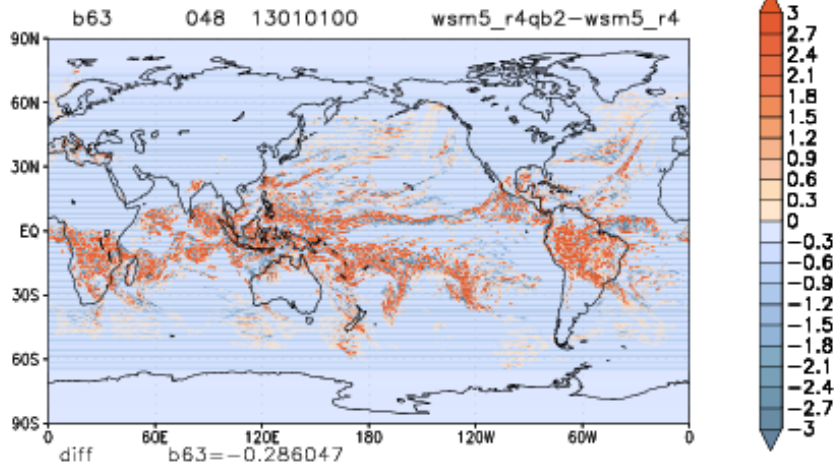
optical depth(ql path, effective radius)

$$\tau = f\left(\frac{\text{particle path} = \int_{ps}^{p^{top}} ql(p) dp}{\text{effective radius}(rq)}\right)$$

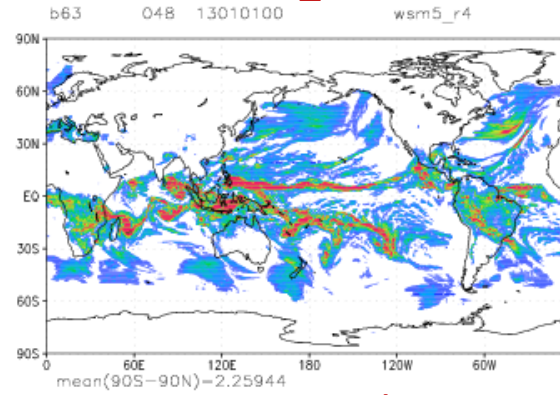
	r2	r4
Cloud water	qc_rad=f(qc+qi+qs+qr, T) rqc=f(T)	qc, rqc =f(qc) from wsm5
Cloud ice	qi_rad=f(qc+qi+qs+qr, T) rqi=f(T)	qi, rqi =f(qi) from wsm5
Rain drop	x	qr, rqr =1000um
snow	x	qs ,rqs =f(qs) from wsm5

# postq ajust impact on precipitation

**Cup**  $\text{diff}(r4qb2-r4) = -0.286 \text{ mm/day}$

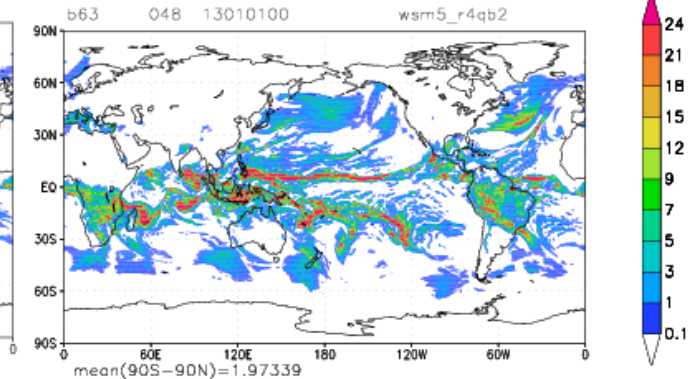


**Wsm5\_r4**



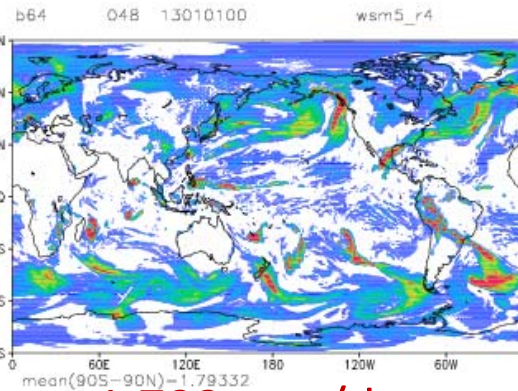
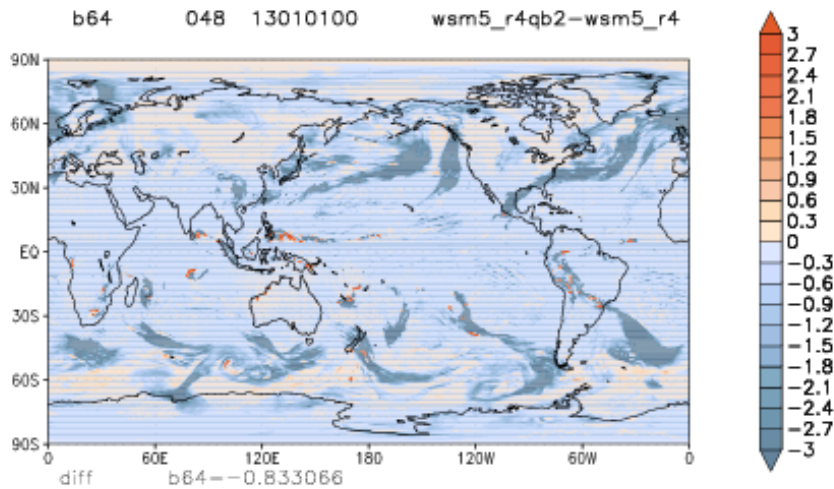
**2.259 mm/day**

**Wsm5\_r4qb2**

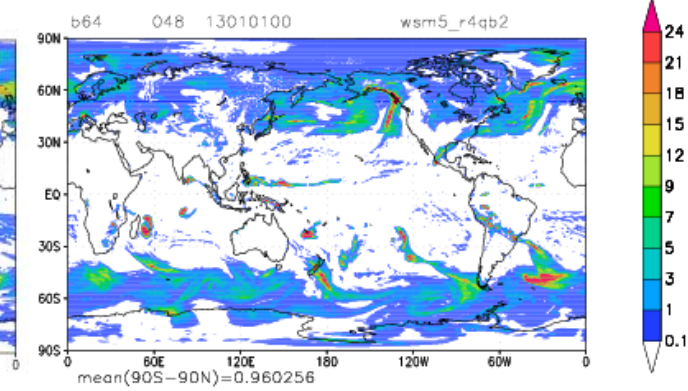


**1.973 mm/day**

**Isp**  $\text{diff}(r4qb2-r4) = -0.833 \text{ mm/day}$



**1.793 mm/day**



**0.960 mm/day**

## 飽和調整策略

當水氣壓超過飽和水氣壓，計算水氣混合比超出飽和混合比之量:  $qv - qvs = sat$

調整第一步，檢查是否有過飽和水氣，若存在過飽和，

檢查是否有**雨滴**存在，若有雨滴存在，計算 $prevp(qv > q_r)$ ，  
但凝結量限制不超過 $sat/2$

調整第二步，檢查過飽和水氣是否被消耗完，若仍存在過飽和，

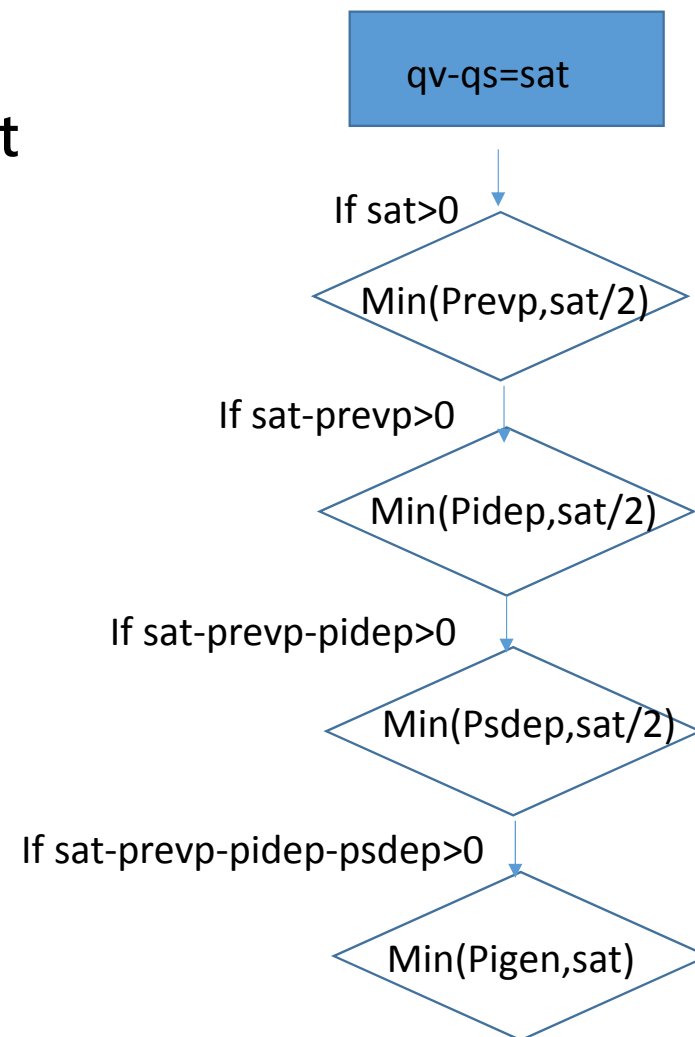
檢查是否有**雲冰**存在，若有雲冰存在，計算 $pidep(qv > q_i)$ ，  
但凝結量限制不超過 $sat/2$

調整第三步，檢查過飽和水氣是否被消耗完，若仍存在過飽和，

檢查是否有**雪**存在，若有雪存在，計算 $psdep(qv > q_s)$ ，  
但凝結量限制不超過 $sat/2$

調整第四步，檢查過飽和水氣是否被消耗完，若仍存在過飽和，

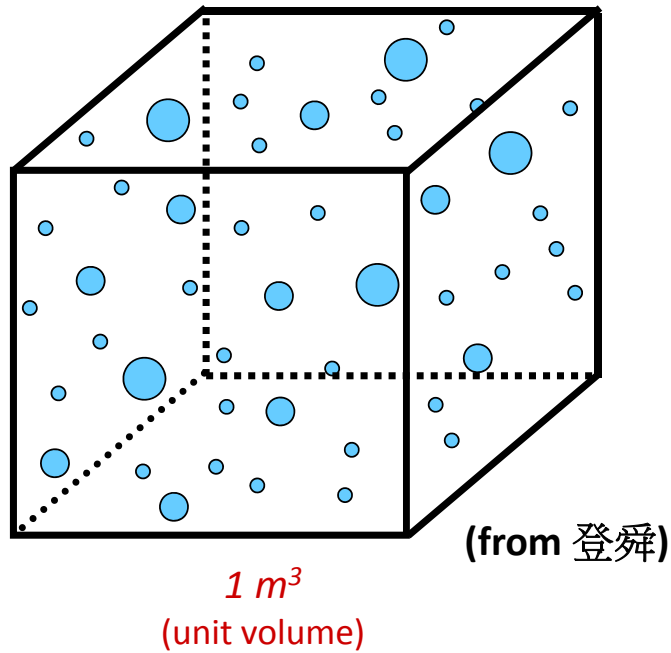
計算**雲冰**產生量， $pigen(qv > q_i)$ ，  
但凝結量限制不超過 $sat$



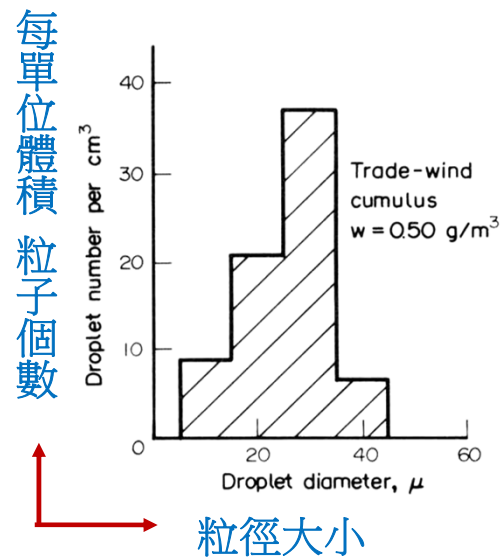


- Nearly all cloud microphysical processes are size dependent

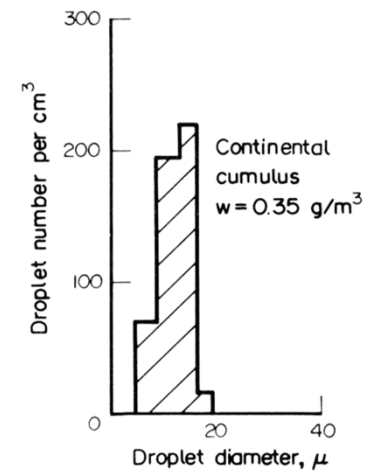
## 粒徑分布(粒徑譜)



Maritime cloud



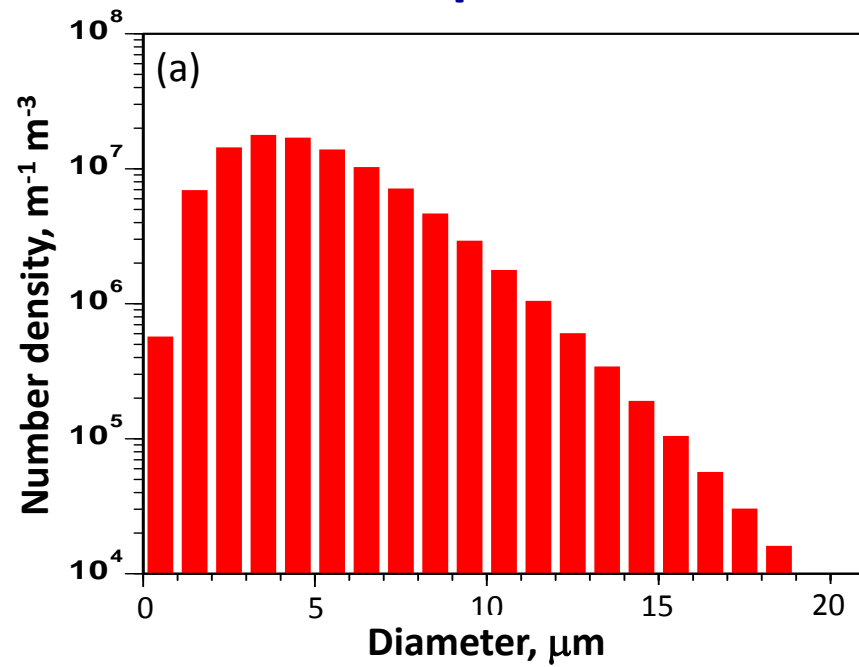
Continental cloud



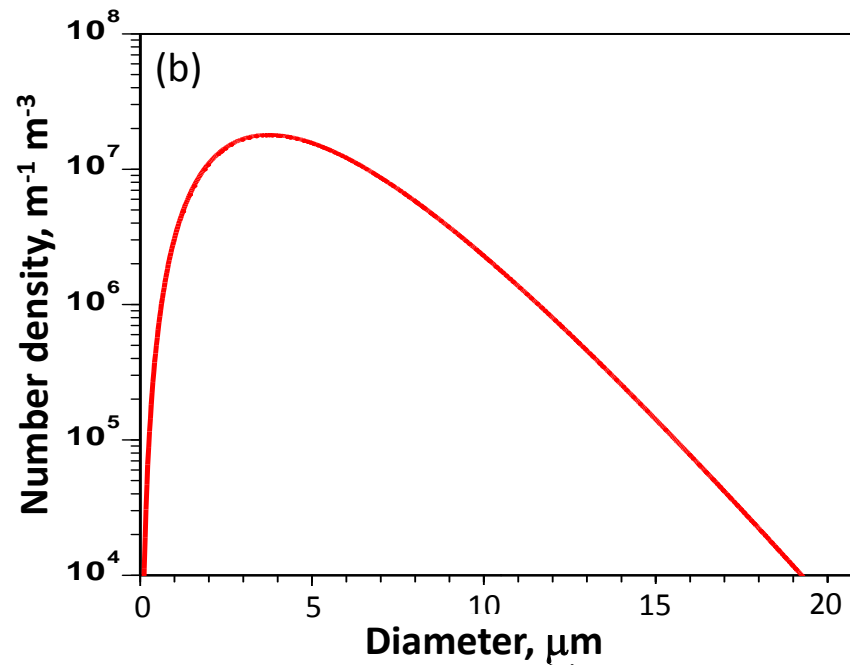
如何處理不同粒徑的雲物理過程

--- different modeling approach

**Spectral**



**Bulk**



(from 陳正平)

## Fall speed in WSM5

- refer to Lin et al. (1983)

Rain, Snow, **cloud ice** 終端速度

$$V_t(\text{rain}) = a_v \times \frac{\Gamma(4 + b_v)}{6} \lambda_r^{b_v} \left( \frac{\rho_0}{\rho} \right)^{0.5}$$

$$V_t(\text{snow}) = a_v \times \frac{\Gamma(4 + b_v)}{6} \lambda_s^{b_v} \left( \frac{\rho_0}{\rho} \right)^{0.5}$$

Parameter		WSM5
Cloud	$a_v$	X
	$b_v$	
Rain	$a_v$	841.9
	$b_v$	0.8
snow	$a_v$	11.72
	$b_v$	0.41
<b>Cloud ice</b>	$V_t(\text{ice}) = 3.29(\rho q_i)^{0.16}$ <i>Heymsfield and Donner(1990)</i>	

# Bulk method Parameterization

## Step 1 Size distribution representation

Marshall–Palmer distribution

$$\text{rain drop } N_r(D) = n_{0r} \exp(-\lambda_r D)$$

$n_{0r}$  為截距參數 =  $8 \times 10^6$

$\lambda_r$  為斜率參數

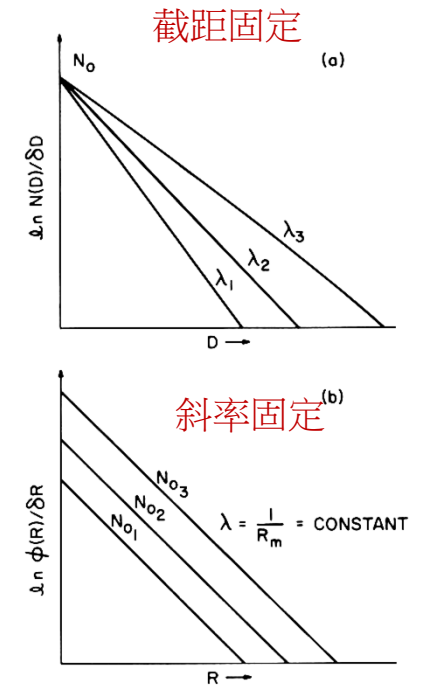
➤ Total number concentration - the 0<sup>th</sup> moments of a size distribution

$$\text{total } N_r = \int N_r(D) dD = \int D^0 N_r(D) dD$$

➤ Mass Mixing ratio - the 3<sup>rd</sup> moments of a size distribution

$$\text{total } q_r = \int q(D) dD = \frac{\pi \rho_r}{6 \rho} \int D^3 N_r(D) dD$$

$$q(D) = \frac{4\pi}{3} \left(\frac{D}{2}\right)^3 N_r(D) \left(\frac{\rho_r}{\rho}\right) \quad \text{for } D \text{ to } D+dD$$



## Bulk method Parameterization step

### Step 2 Get slope parameter $\lambda$

from Mass Mixing ratio - the 3<sup>rd</sup> moments of a size distribution

= $\Gamma(4)$

$$q_r = \int q(D) dD = \frac{\pi \rho_r}{6 \rho} \int D^3 N_r(D) dD = \frac{\pi \rho_r}{6 \rho} \frac{(\int_0^\infty (\lambda_r D)^3 n_{0r} \exp(-\lambda_r D) d(\lambda_r D))}{\lambda_r^4}$$
$$= \frac{\pi \rho_r n_{0r}}{6 \rho \lambda_r^4} \Gamma(4)$$

$$\lambda_r = \left( \frac{\pi \rho_r n_{0r}}{\rho q_r} \right)^{\frac{1}{4}}$$

$q_r$  為預報變數，可反求得到  $\lambda$

### - single moment

增加  $q_r$  (mass mixing ratio) 的預報， $q_r$  同時是粒徑分布函數的一個矩量(3<sup>rd</sup>)