

A modeling study of the influence of multi-scale waves on tropical cyclone formation in June 2004

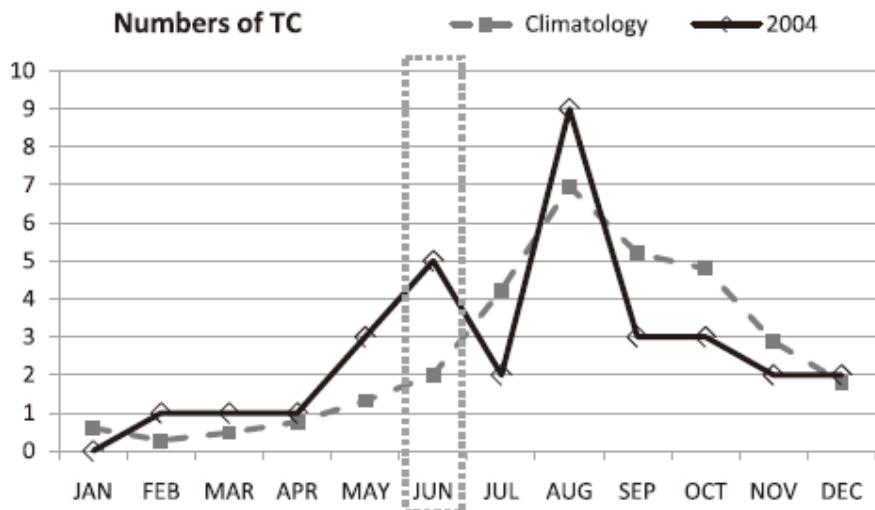
Lin Ching, Chung-Hsiung Sui, Ming-Jen Yang

Department of Atmospheric Sciences, National Taiwan University, Taiwan

TY season: Jul. to Oct.

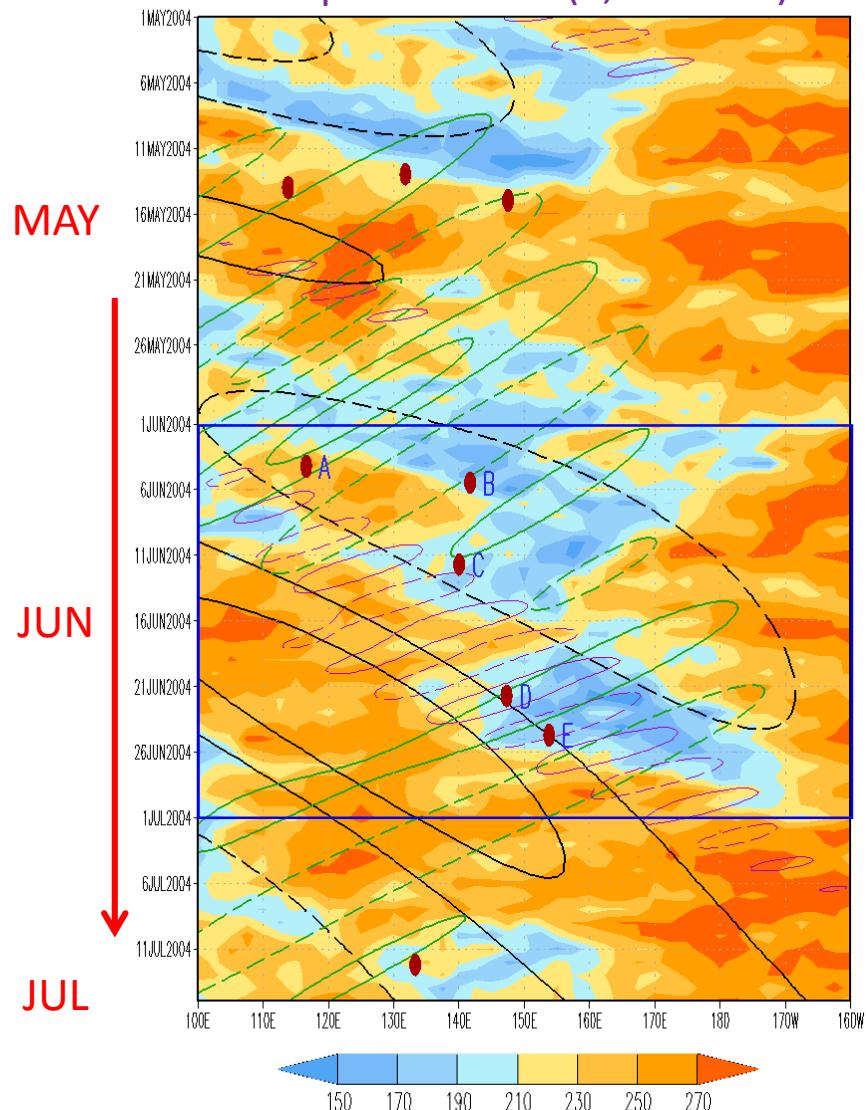
June : transition month

Large-scale conditions are not favorable!

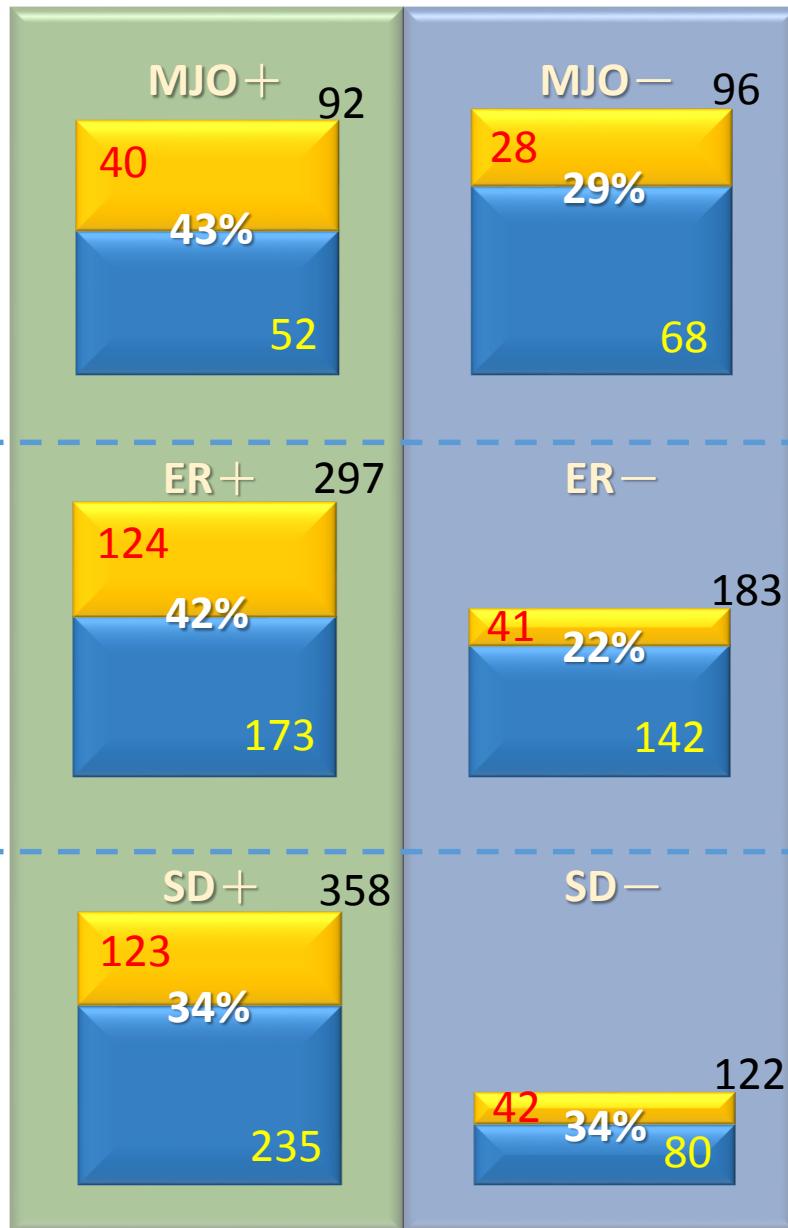


But the record breaking five TCs formed in June 2004

Color: Total OLR (5S-5N)
Black: MJO (OLR, 5S-5N)
Green: ER (VOR, 5N-15N)
Purple: MRGTD (V, 5N-15N)



MJO
(per 100 days)



**Synaptic
scale
disturbance**

TC & Multi-scale waves interaction

Intra-seasonal
to large scale

Bi-weekly to
monthly scale

Synaptic scale

Convection and
non-linear process

WRF experiments design:

Model Setting

Two domain: 36, 12 km

35 vertical layers

Buffer zone: 10 points

Cumulus scheme: Grell 3D ensemble scheme

Microphysics scheme: WSM 6-class scheme

PBL scheme: MYJ PBL scheme

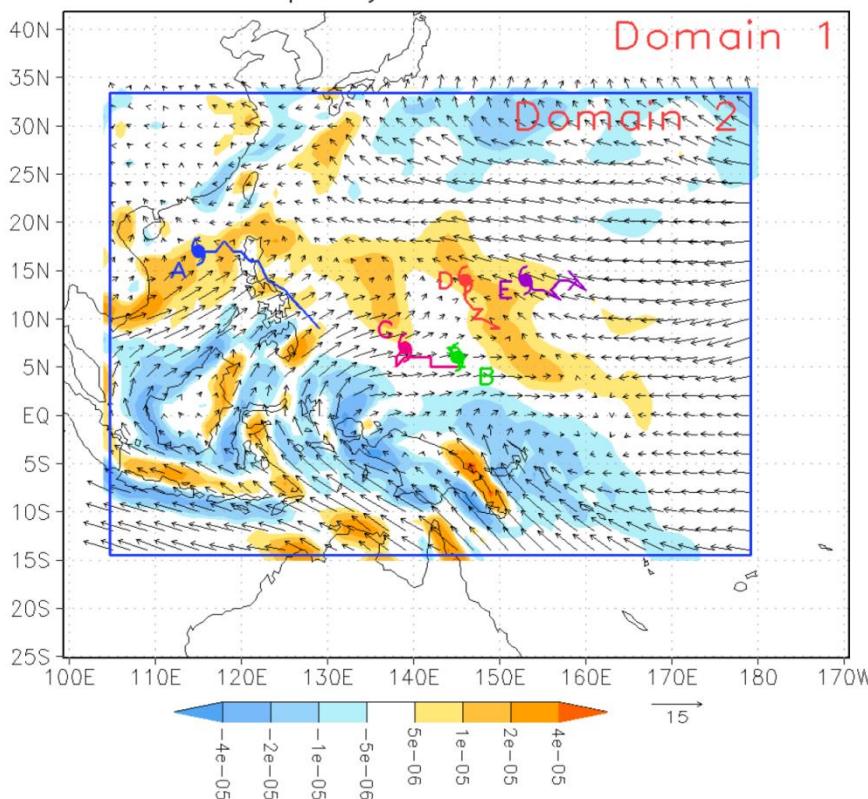
IC & BC: NCEP Global Final analysis data (FNL)

No TC initialization scheme

Space-time filter (*Wheeler and Kiladis, 1999*)

Wave	wave number	period	propagation
MJO	0 – 5	30 – 90 days	eastward
ER	1 – 10	10 – 40 days	westward

low-frequency variations June 2004



Experiment

IC & BC data

CTL

unfiltered NCEP FNL analysis data

noMJO

analysis data excluding MJO signal

noER

analysis data excluding Rossby wave signal

noMJOER

analysis data excluding both MJO and Rossby wave signal

Same setting for all experiments, except the IC & BC data

sensitivity of removed fields

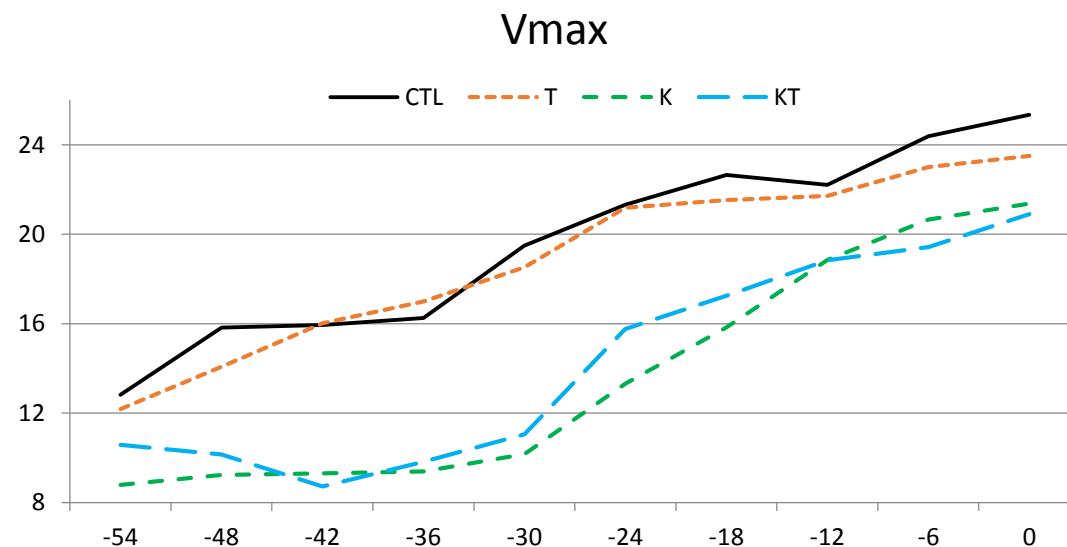
Experiments without wave signal in

T: Thermodynamics fields (H, T, Q)

K: Kinetic fields (U, V)

KT: Kinetic and Thermodynamics fields (U, V, H, T, Q)

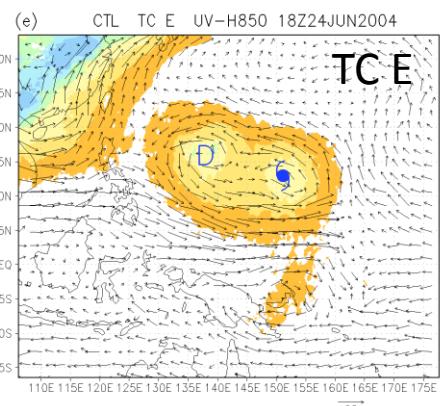
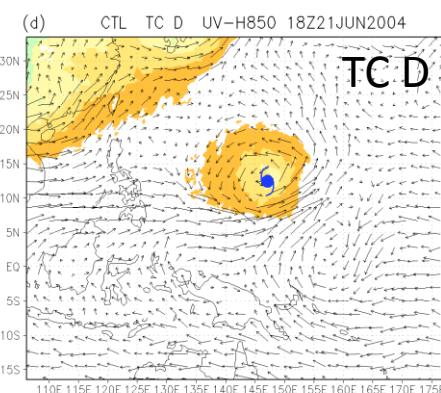
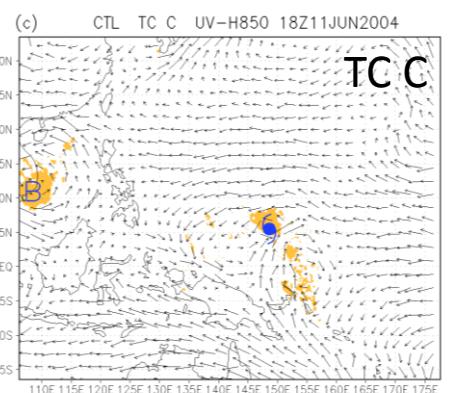
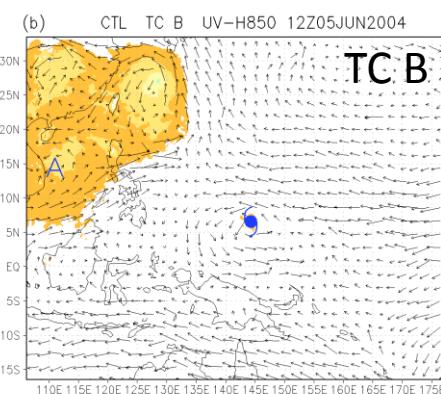
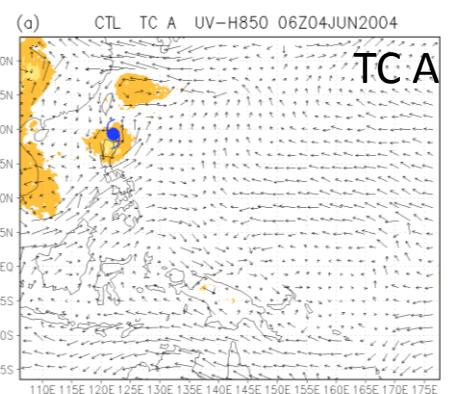
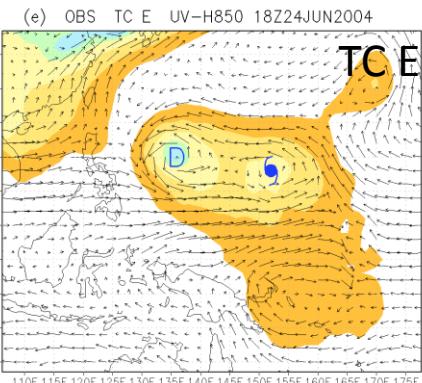
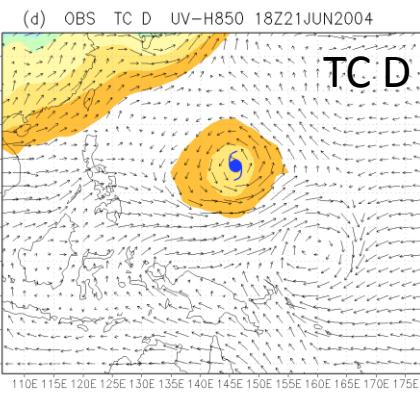
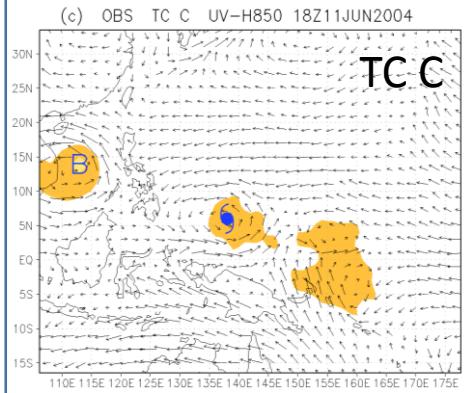
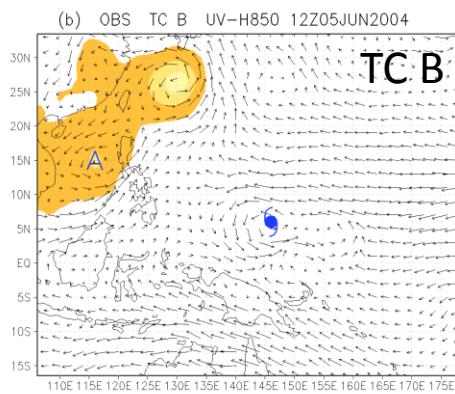
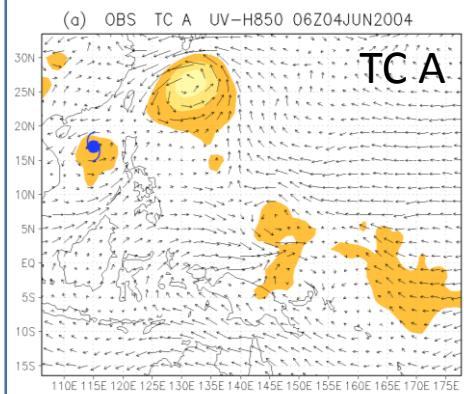
CTL: unfiltered IC & BC



Analysis

HUV850

Simulation



Subtropical Ridge



Tropical Depression (TD):

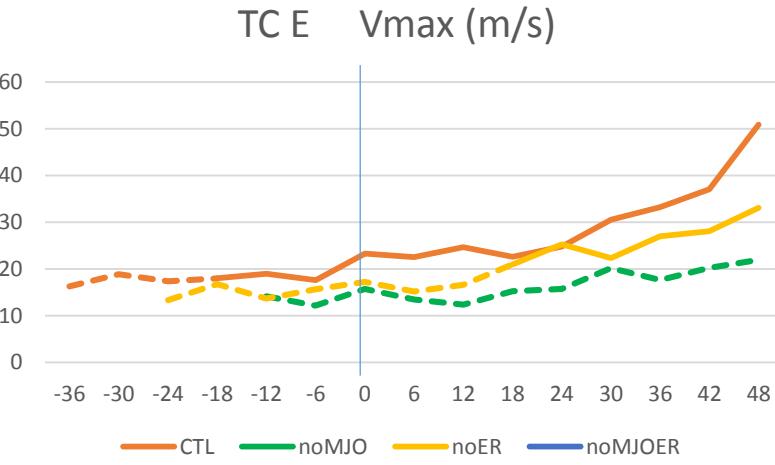
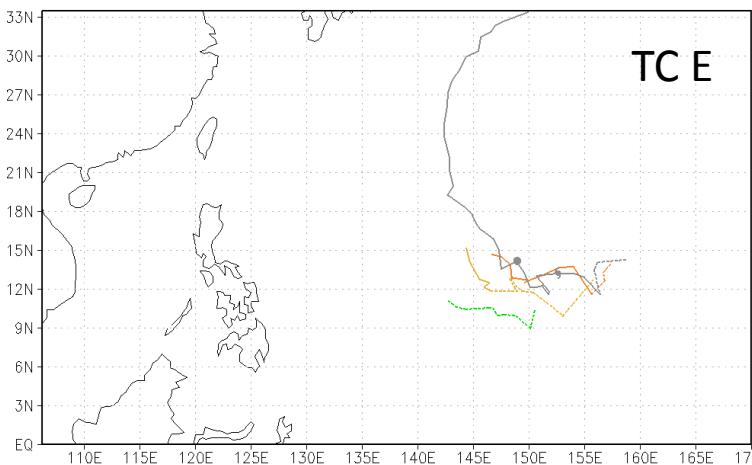
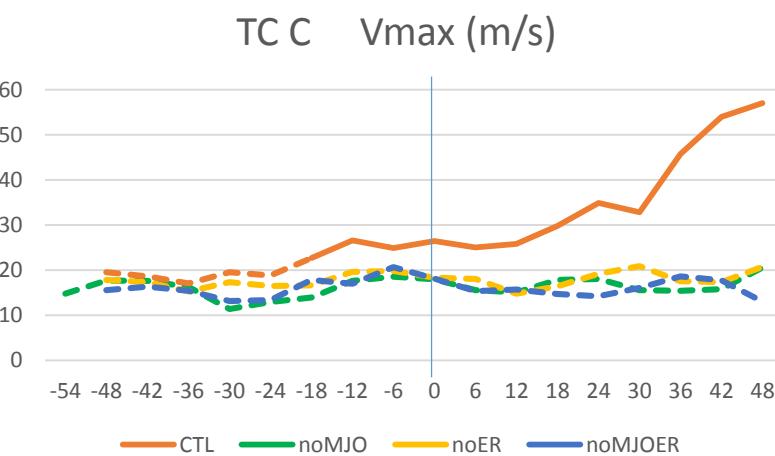
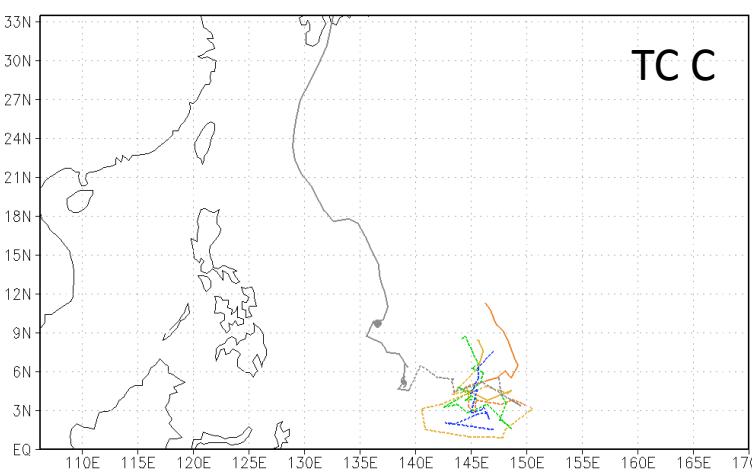
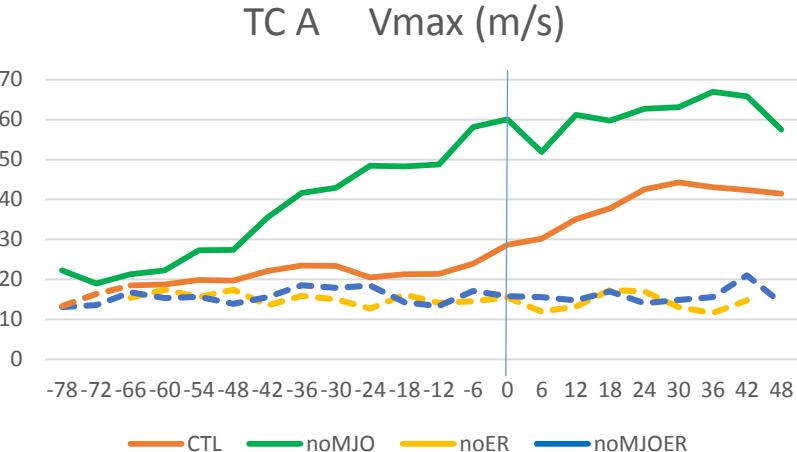
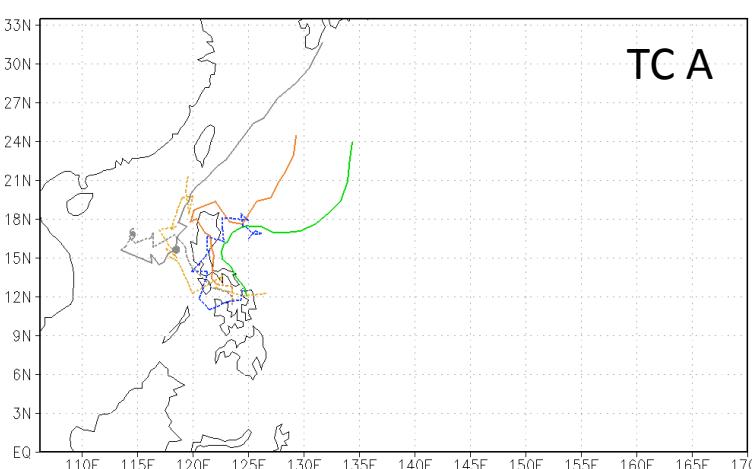
A tropical cyclone in which the maximum sustained wind speed is constantly between 20 and 34 knots. Depressions have a closed circulation.

1. closed circulation
2. $V_{max} > 25 \text{ knots (12.5 m s}^{-1}\text{)}$
3. 14400 km^2 area-mean vorticity $> 5 \times 10^{-5} \text{ s}^{-1}$, and sustain more than 12 hours.

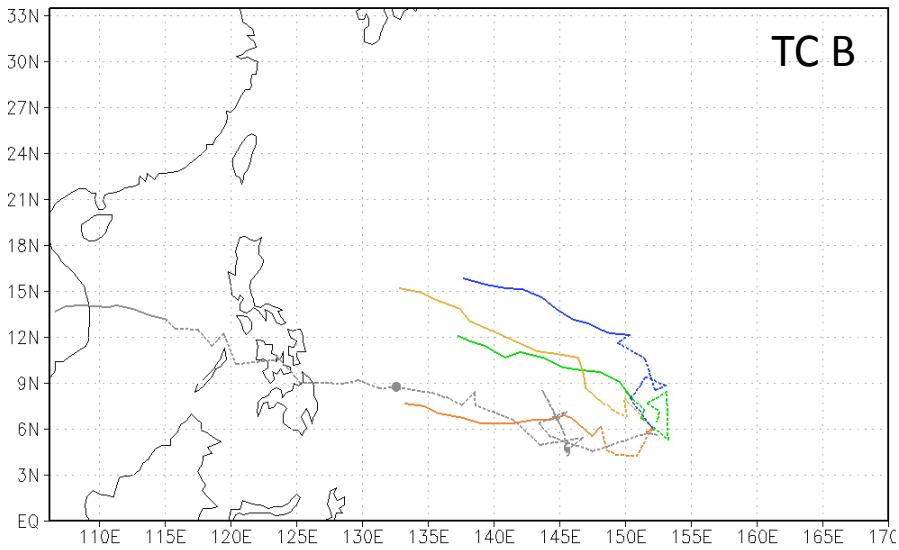
Tropical Storm (TS)

A tropical cyclone in which the maximum sustained wind speed equal or exceed 34 knots.

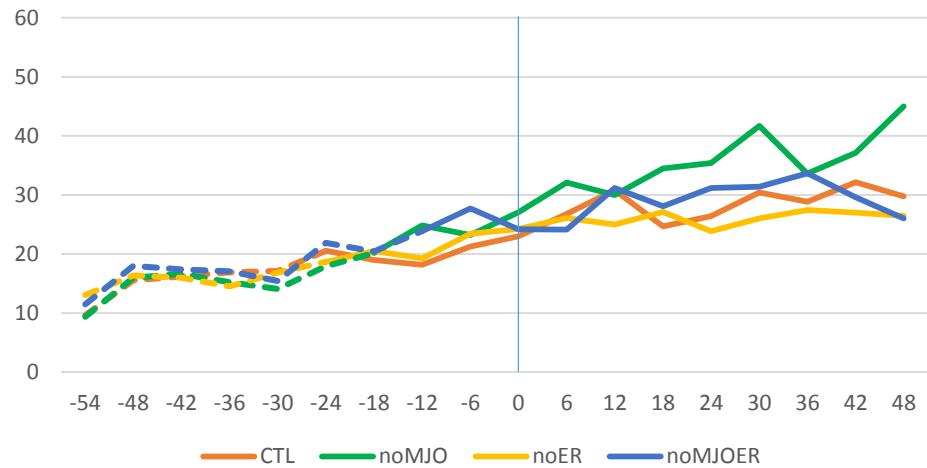
1. closed circulation
2. $V_{max} > 34 \text{ knots (17.5 m s}^{-1}\text{)}$
3. 14400 km^2 area-mean vorticity $> 10^{-4} \text{ s}^{-1}$, and sustain more than 12 hours.



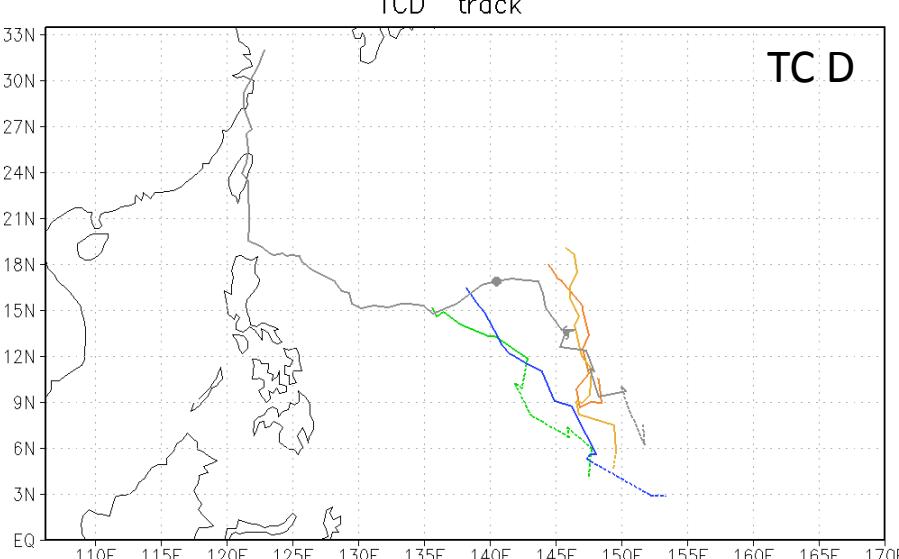
TCB track



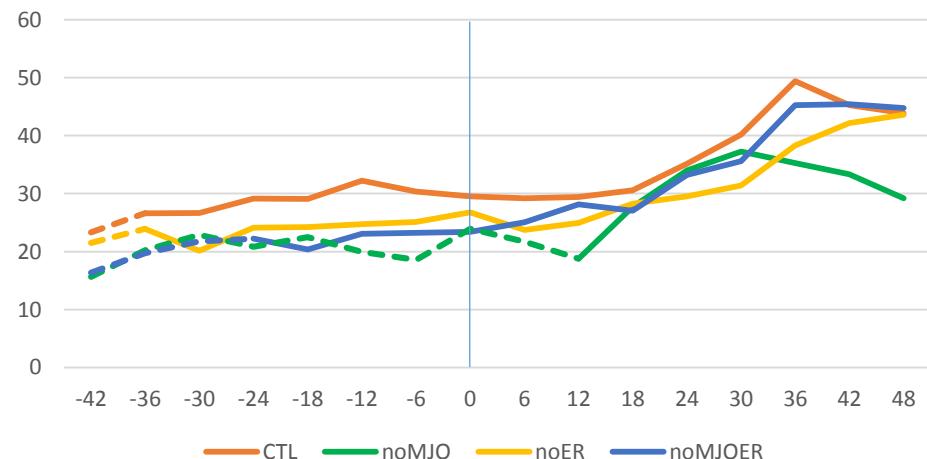
TC B Vmax (m/s)



TCD track



TC D Vmax (m/s)



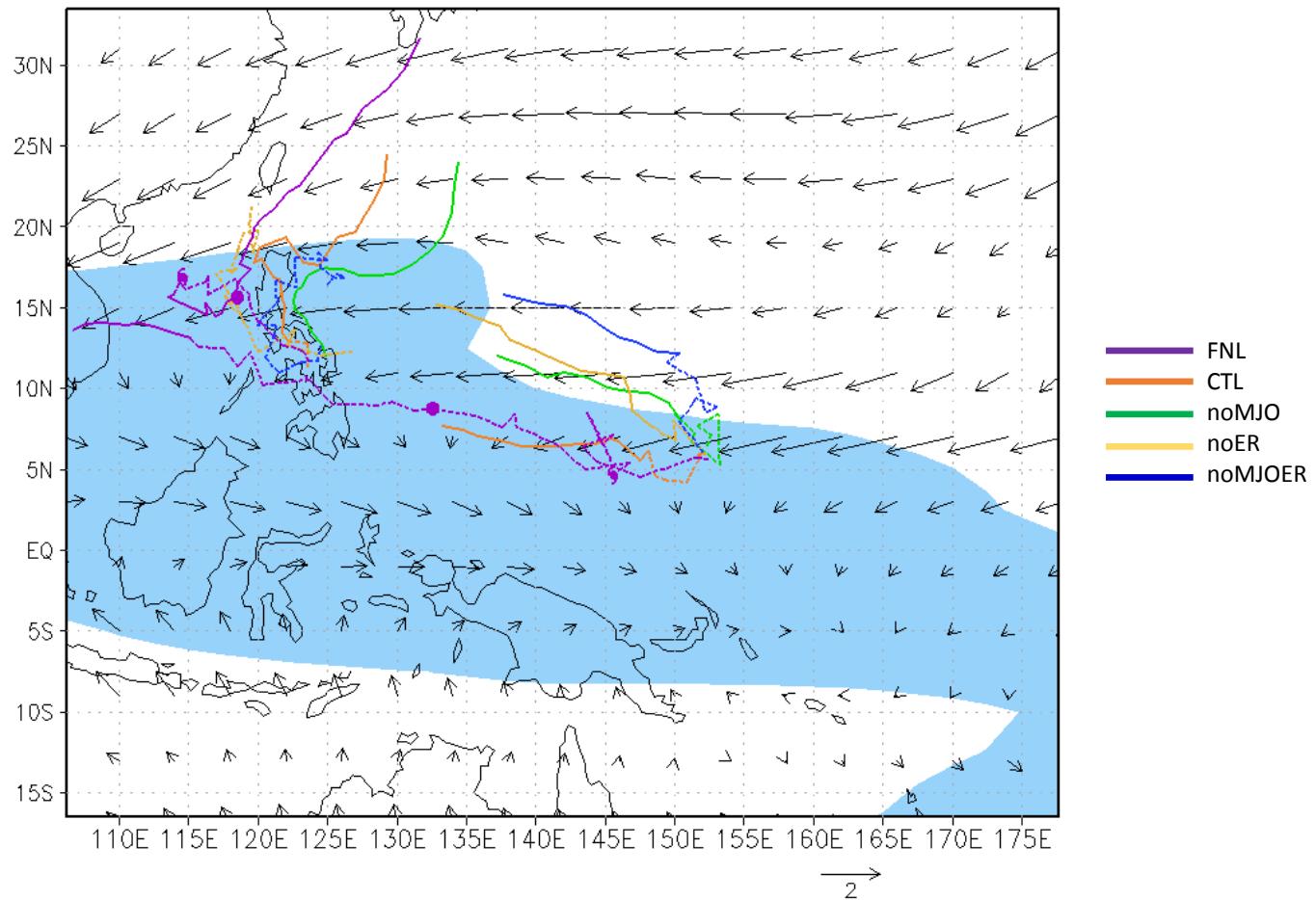
	OBS	CTL	noMJO	noER	noMJO-ER
TC A	TS	TS	TS↑	TCn	TCn
TC B	TS	TS	TS↑	TS	TS
TC C	TS	TS	TCn	TCn	TCn
TC D	TS	TS	TS↓	TS	TS
TC E	TS	TS	TCn	TS↓	disappear

Overall speaking

- Removing LF forcing,
TC tracks change
TC intensity weaken
TS → TD, even disappear

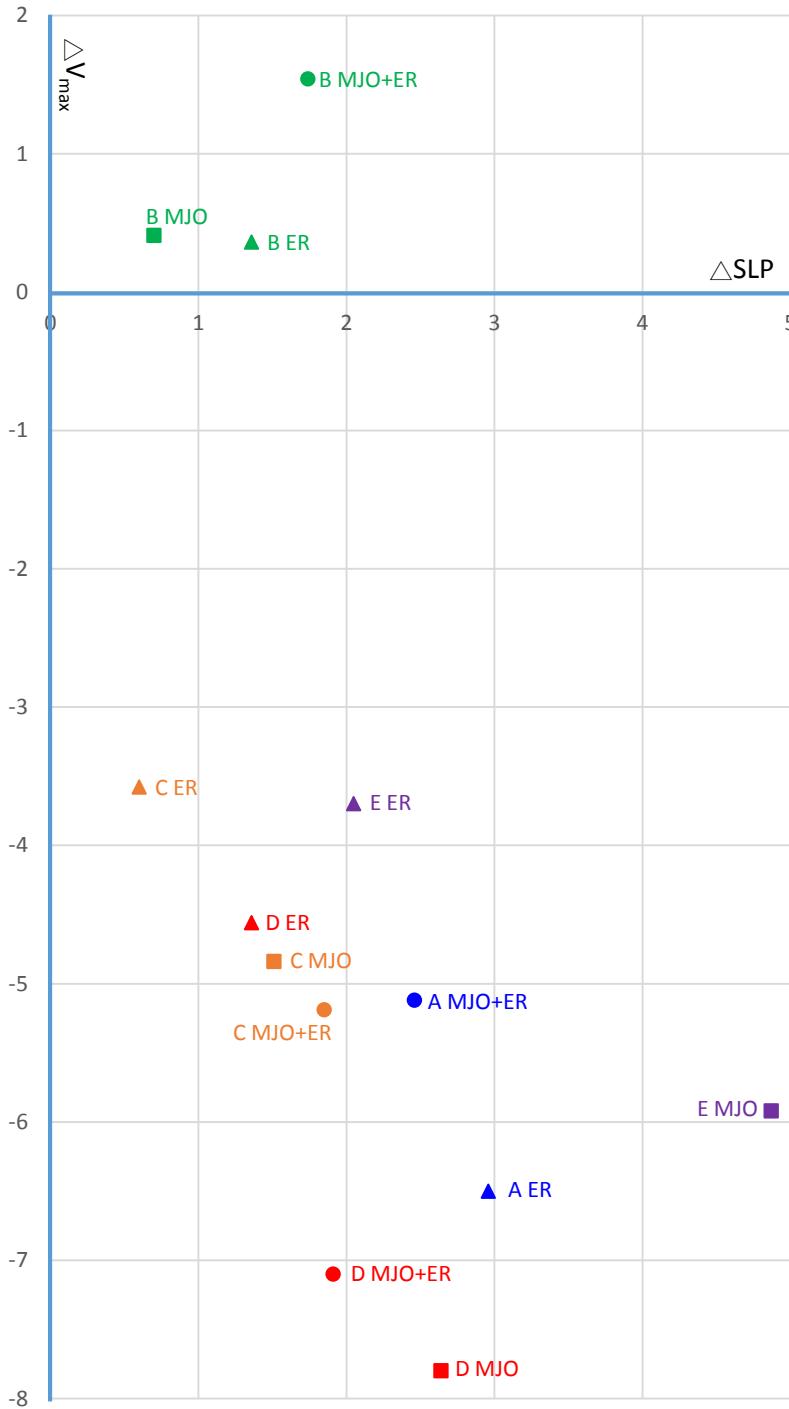
Five days mean MJO signal

5-day-mean (6/3–6/7) OLR UV850 of MJO



Heat flux during developing period (W m^{-2})

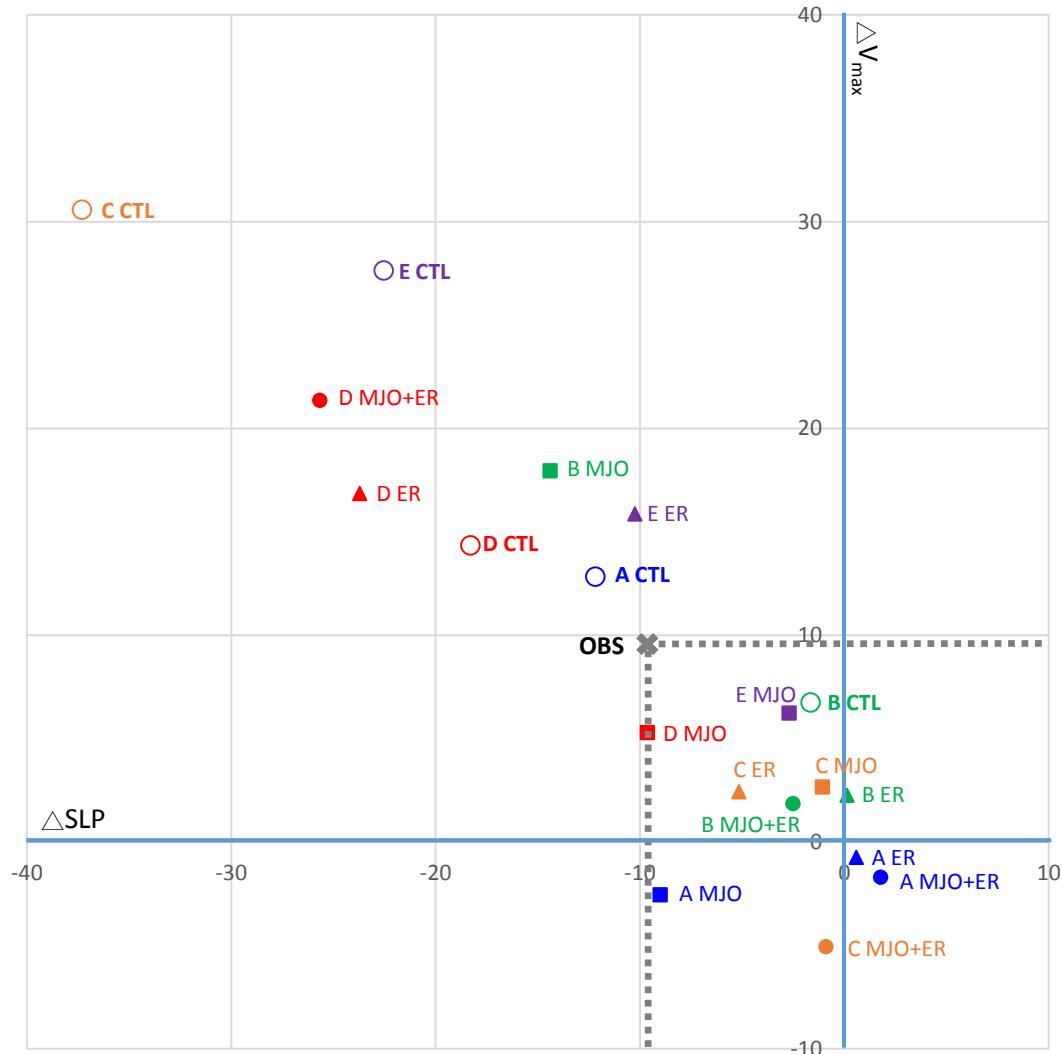
Heat flux	CTL	noMJO	noER	noMJOER
TC A	98.76	159.12	108.83	107.32
TC B	148.1	162.0	143.5	136.4



The differences of SLP (ΔSLP) and V_{max} (ΔV_{max}) of initial disturbance between the CTL and each wave experiment at the warning time $H(0)$ for the five TCs.

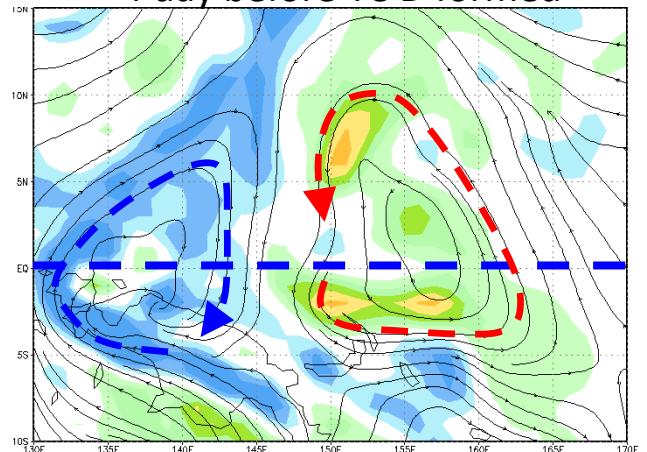
A positive (negative) ΔSLP (ΔV_{max}) denotes a weaker initial disturbance as a result of the missing wave.

An overall strengthening effect by large-scale waves on TC genesis, except TC B.

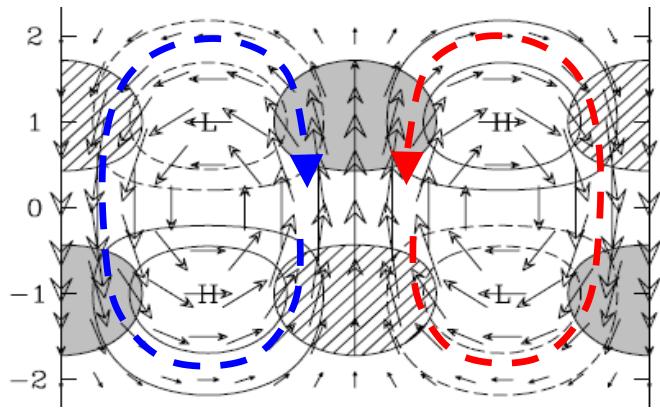


Intensification of observed and simulated TCs in all experiments in two days following the warning time in terms of Δ SLP and ΔV_{max} between the warning time H(0) and 48 hours after the warning time H(48) for the five TCs in June.

4 day before TC D formed

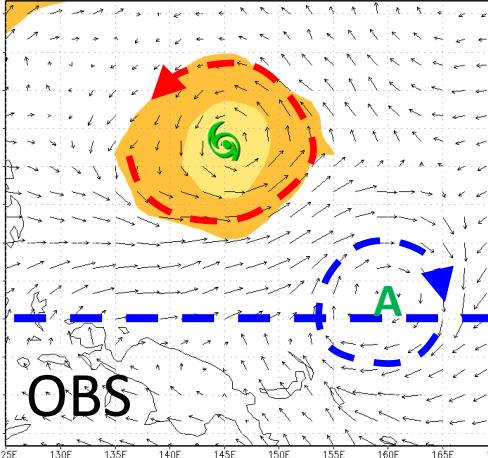


Theoretical MRG wave

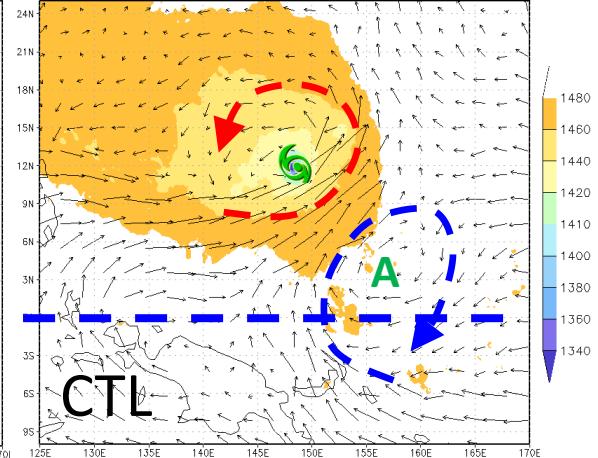


Eq.

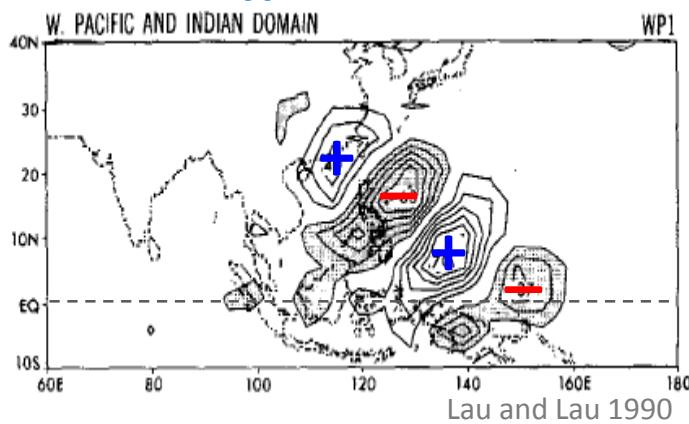
TC D UV-H850 FNL



TC D UV-H850 CTL



TD-type disturbance



Eq.

Moisture Flux Convergence (MFC)

$$(\times 10^{-6} \text{ g kg}^{-1} \text{ s}^{-1}) \quad MFC = -\nabla \cdot (q \vec{V}_h)$$

MFC	TC D
MJO	0.56
ER	-1.08
MRGTD	12.85

CTL

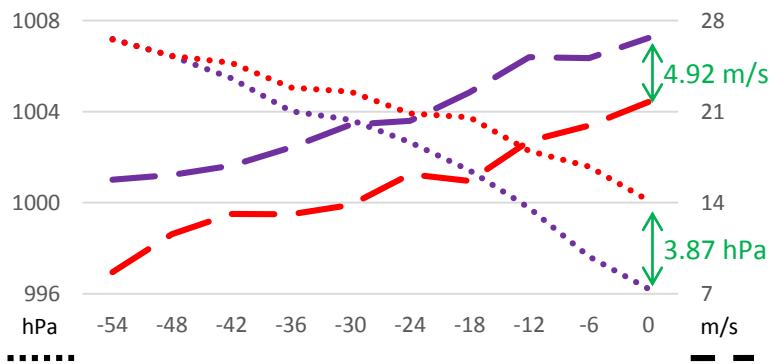
SLP

V_{max} - - -

noMRGTD

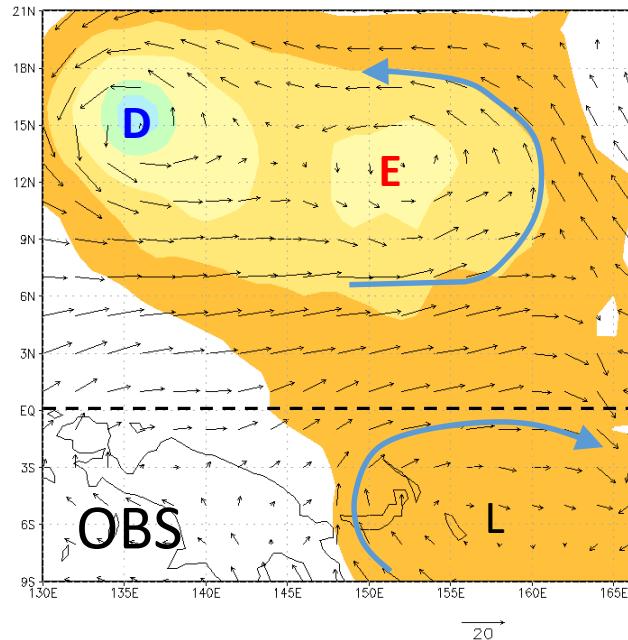
SLP

V_{max} - - -



When TC E formed

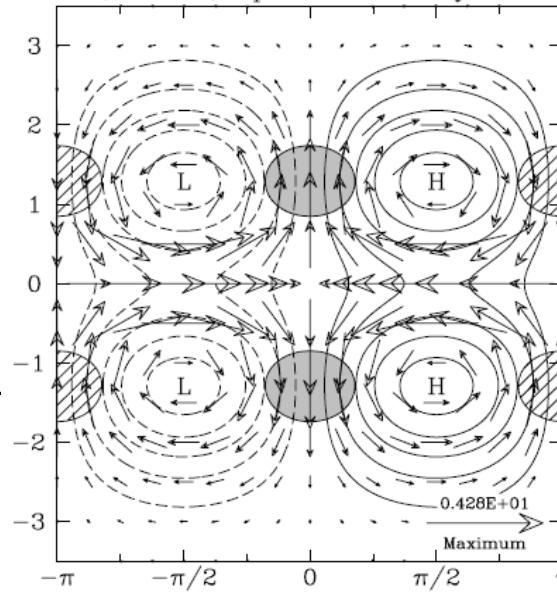
TC E UV-H850 FNL



TC E UV-H850 CTL

Theoretical $n=1$ Rossby wave

$n=1, k^*= -1$, equatorial Rossby



The formation of TC E is mainly contributed by ER wave

mean difference to CTL

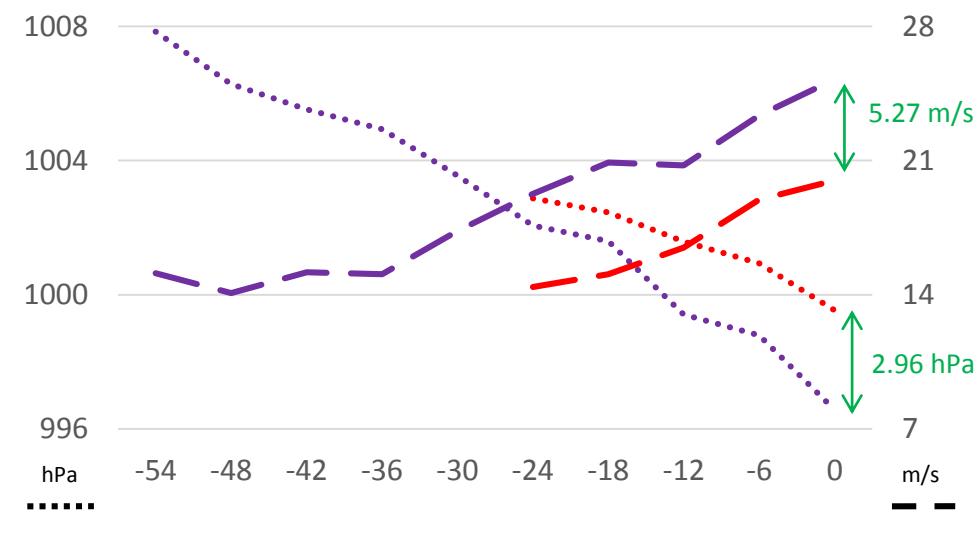
TC E	SLP	V_{max}
no MJO	0.84	-2.56
no ER	1.79	-4.92
no MRGTD	-1.28	-0.39

CTL

noER

SLP

V_{max}



Summary

- The numerical experiments by removing wave signal provide an overall supporting evidence to the observational analysis:
- By conducting a modeling study during this particular and complicated month with active multi-scale wave activities, the five TCs told us some stories:

TC A → the influence of wave is not only on the intensity but also on the track,

TC B → locally-developed system without any influence of wave,

TC C → any wave can dominate the development,

TC D → the contribution of synaptic wave is also important,

TC E → the combined contribution of waves is more pronounced.

Thank you